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Jiang, Yangao

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Radial Growth Response of *Larix gmelinii* to Climate along a Latitudinal Gradient in the Greater Khingan Mountains, Northeastern China

Yangao Jiang 1,2, Junhui Zhang 1,*, Shijie Han 1,*, Zhenju Chen 3, Heikki Setälä 4, Jinghua Yu 1, Xingbo Zheng 1, Yingtao Guo 5 and Yue Gu 1,2

1 Institute of Applied Ecology, Chinese Academy of Sciences, Shenyang 110016, China; jiangyangao-jyg@163.com (Y.J.); yujh@iae.ac.cn (J.Y.); zhxb@iae.ac.cn (X.Z.); guyue740147940@126.com (Y.G.)
2 University of Chinese Academy of Sciences, Beijing 100049, China
3 College of Forestry, Shenyang Agricultural University, Shenyang 110866, China; zhenjuchen@hotmail.com
4 Department of Environmental Sciences, University of Helsinki, Lahti 15140, Finland; heikki.setala@helsinki.fi
5 Key Laboratory of Forest Plant Ecology, Northeast Forestry University, 26 Hexing Road, Harbin 150040, China; gyingtao@163.com

* Correspondence: jhzhang@iae.ac.cn (J.Z.); hansj@iae.ac.cn (S.H.); Tel.: +86-24-8397-0443 (J.Z. & S.H.)

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**Abstract:** To explore how climatic factors influence tree growth within the context of global climate changes, we used a dendroclimatological analysis to understand the response of *Larix gmelinii* to climatic variations along a broad latitudinal gradient from 47.27° to 52.66° N in the Greater Khingan Mountains of Northeastern China. The growth-limiting climate factors and a detailed association between radial growth and climate were identified along the gradient using redundancy analysis (RDA) and standard correlation function analysis over the period 1960–2013. The results showed that temperatures during current June to July represented the most important factor affecting tree radial growth in the study area. Across all studied latitudes, *Larix gmelinii* growth might be decreasing in radial growth due to higher monthly maximum temperature (Tmax) and monthly mean temperatures (Tm) in the current June, especially for the stands at low and middle latitudes. With continued warming, *Larix gmelinii* radial growth at high latitudes (e.g., Mangui (MG) and Mohe (MH)) might be reduced by warmer temperatures in July. In addition, *Larix gmelinii* might be decreasing in radial growth from decreasing precipitation. Our results show that there is a decreasing trend in *Larix gmelinii* radial growth under the observed general increase of temperatures in the Greater Khingan Mountains in recent years.

**Keywords:** boreal forest; dendroecology; tree-ring growth; climate warming

1. Introduction

The stem radial growth of trees in any year integrates the effects of the climate during the previous and current year [1–4]. Tree radial growth is influenced by environmental factors (e.g., temperature, precipitation, and competition) and conditions within the individual tree [1,2]. Tree-ring records are often used to investigate the responses of tree growth to historical climate variations, to predict the effects of future climate change on tree growth, and to understand the spatial and temporal patterns of tree-growth variability of forest ecosystems [5–7]. Therefore, tree-ring records can increase our capacity to predict potential changes in forest structure and composition, as well as provide important information on natural resource management [8–12].
Larix gmelinii (Rupr.) is one of the most economically important species in boreal forests [13]. It is a deciduous needleleaf conifer that is adapted to growing in very harsh climates and is widely distributed over a range from 40° N to 72° N and 110° E to 130° E [13]. It is the dominant tree species in the Greater Khingan Mountains forest ecosystem in northern China [14,15].

Modeling studies predict that L. gmelinii growth will decrease rapidly in China due to the predicted increases in temperature [16–19] and decreases in precipitation [17]. However, empirical/observational evidence of such declines in growth at the regional scale is lacking in this region. Only a few investigations on L. gmelinii radial growth response to climate variation have been conducted at a small number of distinct locations in the northern part of the Greater Khingan Mountains during the past few decades [20–22]. The aims of this paper are (1) to identify the spatial pattern of the radial growth-climate relationship of Larix gmelinii and (2) to identify the dominant climatic factors affecting growth over a large latitudinal range (47.27° to 52.66° N) in the western Greater Khingan Mountains, China.

2. Materials and Methods

2.1. Study Area

The study was conducted along a latitudinal gradient in eastern Inner Mongolia and Heilongjiang Province of Northeastern China. Tree-ring samples were collected at seven sites: Mohe (MH), Mangui (MG), Moerdaoga (MEDG), Yitulihe (YTLH), Wuerqihan (WEQH), Boketu (BKT), and Aershan (AES), ranging from 120.30° to 122.31° E and 47.27° to 52.66° N (Figure 1). All of the selected stands of Larix gmelinii were naturally established, with no evidence of recent disturbance. To minimize the effects of local terrain and microenvironmental factors on tree growth, the topography of all selected plots was uniform and flat, with an elevation range of 908 m to 1018 m and a southern aspect with the slope ranging from 0 to 20 degrees (Table 1). Further, all sites except AES were located at the peaks of the local mountains. The sampling sites cover almost the entire native latitudinal distribution range of Larix gmelinii in China. The climate in the study region is continental monsoon characterized by a windy spring, a short, warm, and humid summer, and a long, dry, and cold winter. The mean annual precipitation (1960–2013) of the study sites ranged between 415 and 488 mm (Figure 2), with the period between April and September accounting for ~88%–90% of the annual precipitation [23]. The mean annual temperature ranged from −5.8°C to −1.8°C [23]. Soil type varies among the seven sites, being mostly brown coniferous forest soils, chernozem soil, and grey forest soil, with a depth of 8 to 100 cm (Table 1).

Figure 1. Map showing the locations in China where Larix gmelinii tree-ring samples were collected (red-filled circles). Abbreviations for our sampling sites are AES (Aershan), BKT (Boketu), WEQH (Wuerqihan), YTLH (Yitulihe), MEDG (Moerdaoga), MG (Mangui), and MH (Mohe).
were fixed in wooden holders designed to hold the cores firmly, and they were then carefully polished. In the laboratory, the tree-ring samples were air-dried, after which they were cross-dated to remove the biological growth related to the tree’s age. The cross-dated tree-ring data were detrended by fitting negative exponential or linear curves [25]. Cross-dating was verified using COFECHA software [24], which calculates correlation coefficients between individual tree-ring series as a way to identify absent or false rings.

Several commonly used descriptive statistics were calculated for comparisons among the chronologies. The standard deviation (SD) and the mean sensitivity (MS) were used to assess the high-frequency variations of the tree-ring chronologies [1], the first-order autocorrelation (AC1) was calculated using CANOCO software [28], and the percentage of the total variance explained by the first two RDA components (from 1960 to 2013) was extracted from the $\mathcal{X}$ covariance matrix. Analyses were then carried out using the program RDA [28]. Cross-dating was also conducted at a 1-year period using crossdating software [24], which is used to analyze the cross-dating results of the tree-ring series. The cross-dating results were then used to extract the chronologies for each site. The chronologies were then averaged to generate a standard (STD) chronology using the program ARSTAN [26]. In total, seven residual ring-width chronologies of Larix gmelinii were constructed and used in the subsequent analysis (Figure 3). Several commonly used descriptive statistics were calculated for comparisons among the chronologies. The standard deviation (SD) and the mean sensitivity (MS) were used to assess the high-frequency variations of the tree-ring chronologies [1], the first-order autocorrelation (AC1) was calculated using CANOCO software [28], and the percentage of the total variance explained by the first two RDA components (from 1960 to 2013) was extracted from the $\mathcal{X}$ covariance matrix. Analyses were then carried out using the program RDA [28]. Cross-dating was also conducted at a 1-year period using crossdating software [24], which is used to analyze the cross-dating results of the tree-ring series. The cross-dating results were then used to extract the chronologies for each site. The chronologies were then averaged to generate a standard (STD) chronology using the program ARSTAN [26].

2.2. Tree-Ring Sampling and Chronology Development

Each site consisted of three stands. Twenty-one forested stands were sampled in 2013 and 2014. Two tree-ring increment cores per tree were extracted from the trunk at breast height. In total, 157 trees (285 cores) were sampled. In the laboratory, the tree-ring samples were air-dried, after which they were fixed in wooden holders designed to hold the cores firmly, and they were then carefully polished with successively finer grits of sandpaper. The tree rings were visually cross-dated under a binocular microscope, and the ring widths were carefully measured using a LINTAB5 measuring system with a precision of 0.001 mm. Cross-dating was verified using COFECHA software [24], which calculates correlation coefficients between individual tree-ring series as a way to identify absent or false rings. The cross-dated tree-ring data were detrended to remove the biological growth related to the tree’s age and any other non-climate-related growth variation by fitting negative exponential or linear curves [25]. In a few cases, a cubic spline with a 50% frequency response cutoff equal to two thirds of the series length was utilized when anomalous low-frequency growth trends occurred. These index series were then averaged to generate a standard (STD) chronology using the program ARSTAN [26]. In total, seven residual ring-width chronologies of Larix gmelinii were constructed and used in the subsequent analysis (Figure 3).

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calculated to assess relationships with previous growth, and the signal-to-noise ratio (SNR) was calculated to evaluate the signal strength of the site chronologies [1]. The expressed population signal (EPS) was applied to measure the similarity between the finite-sample chronology and the theoretical population chronology based on an infinite number of trees. The threshold value for accepting an EPS value was set to 0.85 after Wigley et al. (1984) [27]. The mean series inter-correlation (RBAR) was used to evaluate the signal strength of each chronology, which is the mean correlation coefficient among the tree-ring series [22].

![Figure 3. Ring-width residual chronologies for the period 1960–2013 constructed for Larix gmelinii along the latitudinal gradient from 47.27° to 52.66° N at the seven sites (panels) in the Greater Khingan Mountains of the western Chinese boreal forest. The dashed lines indicate the numbers of tree-ring series.]

2.3. Climate Data

There are only a few weather stations in the research area. Consequently, the climate observations for dendroclimatological investigations were insufficient. The climate data of research sites for analysis of growth-climate response were interpolated values on records that were randomly selected from 120 of 164 climate stations (120 stations across northeast China from Chinese Meteorological Data Sharing Service System [28] and 44 stations in Russia from Global Historical Climatology Network [29]).

The method of this interpolation is to fit a partial thin-plate smoothing spline model with location and elevation to ground-based observations for estimating a trend surface and then a simple kriging procedure is applied to the residuals for trend surface correction [30,31]. To evaluate the accuracy of interpolated values, a regression analysis was performed between interpolated and measured temperatures and precipitations at the remaining 44 climate stations. There was a good agreement with $R^2 = 0.99$.

Six climate variables were used for the dendroclimatological analyses, including monthly mean temperature (Tm), monthly minimum temperature (Tmin), monthly maximum temperature (Tmax), monthly total precipitation (Prec), annual mean temperature (Ta), and annual total precipitation (Prec).

2.4. Climate–Growth Relationship Analysis

The correlation analysis for climate and tree growth was performed using the response functions over a 17-month period (from June of the previous year to October of the current year of growth) [1]. The interval parameter was set to 1960–2013, which covered all of the tree-ring series. DENDROCLIM 2002 was used to carry out the statistical calculations [32].
2.5. Definition of Regional Patterns

Redundancy analysis (RDA) was used to analyze the regional responsive patterns of *Larix gmelinii* radial growth to environmental factors over the period 1960–2013. RDA is a multivariate "direct" gradient analysis that constrains its two tables of variables X and Y to represent linear combinations of supplied environmental factors [33]. RDA is a non-symmetric method in which the components extracted from X are correlated as much as possible with the variables of Y, and then the components of Y are extracted with the aim of correlating them as much as possible with the components extracted from X. The correlation between abiotic and biotic variables is revealed by the cosine of the angle between two vectors in RDA biplots. Vectors pointing in the same direction show a high positive correlation, vectors pointing in opposite directions indicate a high negative correlation, and vectors crossing at right angles reveal a near-zero correlation [34]. The longer the vector, the more important the environmental parameter [33]. Each tree-ring chronological series within a site was treated as a sample, and each monthly response function was treated as a descriptor in this analysis [32]. For the statistical quantification of the similarity of the dendrochronological series, we calculated matrix correlation coefficients for the common period 1960–2013. The RDA was carried out for the seven chronologies pooled and the 42 climate variables over the common interval (1960–2013) using the program CANOCO (Version 4.5) [35].

3. Results

3.1. Chronology Statistics

Table 2 shows the dendrochronological characteristics of *Larix gmelinii* ring-width chronologies. The mean ring width ranged between 0.77 and 3.4 mm. The mean sensitivity (MS) and standard deviation (SD) values ranged from 0.221 to 0.308 and from 0.143 to 0.321, respectively, with the highest values at the middle latitudes (i.e., from WEQH to MEDG). For the analysis period of 1960–2013, the EPS of all the chronologies was higher than 0.85. The high inter-series correlation, with values ranging between 0.504 and 0.665, suggests that the chronologies recorded sufficient environmental signals.

<table>
<thead>
<tr>
<th>Site code</th>
<th>MH</th>
<th>MG</th>
<th>MEDG</th>
<th>YTLH</th>
<th>WEQH</th>
<th>BKT</th>
<th>AES</th>
</tr>
</thead>
<tbody>
<tr>
<td>SD</td>
<td>0.178</td>
<td>0.275</td>
<td>0.321</td>
<td>0.229</td>
<td>0.278</td>
<td>0.145</td>
<td>0.143</td>
</tr>
<tr>
<td>MS</td>
<td>0.228</td>
<td>0.221</td>
<td>0.237</td>
<td>0.308</td>
<td>0.253</td>
<td>0.286</td>
<td>0.249</td>
</tr>
<tr>
<td>AC1</td>
<td>0.528</td>
<td>0.590</td>
<td>0.580</td>
<td>0.680</td>
<td>0.610</td>
<td>0.460</td>
<td>0.790</td>
</tr>
<tr>
<td>RBAR</td>
<td>0.665</td>
<td>0.593</td>
<td>0.504</td>
<td>0.660</td>
<td>0.504</td>
<td>0.504</td>
<td>0.582</td>
</tr>
<tr>
<td>EPS</td>
<td>0.965</td>
<td>0.966</td>
<td>0.958</td>
<td>0.983</td>
<td>0.952</td>
<td>0.969</td>
<td>0.958</td>
</tr>
<tr>
<td>PC1</td>
<td>0.369</td>
<td>0.410</td>
<td>0.390</td>
<td>0.628</td>
<td>0.280</td>
<td>0.460</td>
<td>0.730</td>
</tr>
<tr>
<td>SNR</td>
<td>27.355</td>
<td>28.646</td>
<td>22.785</td>
<td>56.848</td>
<td>20.012</td>
<td>30.782</td>
<td>23.089</td>
</tr>
<tr>
<td>MRW (mm)</td>
<td>0.77 ± 0.2</td>
<td>1.13 ± 0.4</td>
<td>0.67 ± 0.2</td>
<td>2.03 ± 0.5</td>
<td>0.68 ± 0.4</td>
<td>2.09 ± 0.5</td>
<td>3.4 ± 0.4</td>
</tr>
<tr>
<td>No. of radii (trees)</td>
<td>41 (22)</td>
<td>43 (28)</td>
<td>41 (26)</td>
<td>40 (27)</td>
<td>37 (25)</td>
<td>40 (32)</td>
<td>39 (25)</td>
</tr>
</tbody>
</table>

SD, standard deviation; MS, mean sensitivity; AC1, first-order autocorrelation; RBAR, mean inter-series correlation; EPS, expressed population signal; PC1, percent variance explained by the first principal component; SNR, signal-to-noise ratio calculated for the standardized chronologies prior to autoregressive modeling; MRW, mean ring width.

3.2. Radial Growth–Climate Associations

Correlation coefficients between the tree-ring growth response of *Larix gmelinii* and the climatic variables were calculated to determine the significance of the climatic factors and the timing (Figure 4). In general, air temperature had a greater influence than the other climatic variations. Radial growth, as a whole, was negatively correlated with monthly mean temperatures (Tm) and monthly maximum temperatures (Tmax) in the current June at all latitudes, with significant correlations were found for
stands at middle and low latitudes. *Larix gmelinii* radial growth across the latitudes was positively correlated with Tm and monthly minimum temperatures (Tmin) during current April to May, with significant correlations observed for AES, BKT and WEQH in May. In addition, radial growth was negatively correlated with current Tm and Tmin in February and March at most latitudes. *Larix gmelinii* radial growth across all latitudes was positively and negatively correlated with Tmin and Tmax, respectively, of the previous June, with significant correlations observed for AES, YTLH and MG/WEQH. In addition to the common growth response to climate across the gradient, site-specific growth responses to climate were also observed (Figure 4). For example, the radial growth of *Larix gmelinii* at AES was negatively influenced by the Tm of the previous December. In contrast, the WEQH and MEDG stands showed significant positive correlations with the September Tm of the previous year. Radial growth at MH and MG was significantly negatively correlated with Tm and Tmin in the current-year July. In addition, the radial growth at AES was also significantly positively and negatively correlated with Tmin in the current July and previous August, respectively.

Figure 4. Bootstrapped correlation coefficients between *Larix gmelinii* residual tree-ring chronologies and the monthly climate variables (monthly maximum, mean, and minimum temperatures, monthly total precipitation, and monthly drought code) for the period 1960–2013 along the latitudinal gradient from 47.27° to 52.66° N in the boreal forest of the western Greater Khingan Mountains in Northeastern China. The black dots indicate significant correlations at the p < 0.05 level.
Precipitation also had a significant influence on the radial growth of *Larix gmelinii* (Figure 4). At almost all latitudes, the positive correlation was observed between radial growth and precipitation during March to August of the current year and summer to autumn of the previous year. However, the correlation was negative for the winter of the previous year. The radial growth was significantly and positively correlated with the current year’s PDSI in the current June–July at WEQH, and a significantly negative correlation was observed in the current June at BKT.

### 3.3. Spatial Patterns of Climate–Growth Responses

In terms of climate, the dendrochronological series could be separated into three groups in the RDA (Figure 5). The first group of series correlated positively with \( T_a \); this group comprised the sites located south of 49°N (low-latitude region). The second group of the dendrochronological series (MG and MH), at latitudes higher than 52° N and situated in the high latitudes of the region, showed a significantly negative relationship with \( T_a \). For the stands located in the middle latitudes of the region (from 49° to 51° N, i.e., WEQH, YTLH, and MEDG), the importance of local factors (e.g., soil N content, slope, soil depth) increased.

![Redundancy analysis (RDA) calculated from the seven residual chronologies, local factors, and the monthly climate parameters for the period 1960–2013.](image)

**Figure 5.** Redundancy analysis (RDA) calculated from the seven residual chronologies, local factors, and the monthly climate parameters for the period 1960–2013. Significant \( p < 0.05 \) environmental factors are indicated by vectors (blue arrows); the longer the vector, the more important the climate parameter. The correlation between the variables is illustrated by the cosine of the angle between two vectors. Vectors at sharp angles indicate a positive correlation, vectors at obtuse angles indicate a negative correlation, and vectors crossing at right angles indicate a near-zero correlation [33]. \( T_a = \) annual mean temperature, \( \text{Prec} = \) annual total precipitation, \( \text{N} = \) soil organic nitrogen content. Slope = the slope of the sampling point, Soil depth = the depth of the soil layer at the sampling point, Lat = the latitude of the sampling point. MH, MG, MEDG, YTLH, WEQH, BKT, and AES represent the seven sampling sites (i.e., Mohe, Mangui, Moerdaoga, Yitulihe, Wuerqihan, Boketu, and Aershan, respectively).

To verify whether the spatial patterns of the radial growth response to climate varied with latitude, a correlation coefficient matrix among the studied chronologies was calculated and analyzed (Table 3). The results revealed a high correlation in the dendrochronological series between neighboring latitudes. The highest correlation was that between MEDG and WEQH, while the lowest correlation was that
between MH and AES (Table 3). Furthermore, we found two relatively high correlations; one was the series from latitudes lower than 49° N; the other was the series from latitudes higher than 52° N. The RDA (Figure 5) and correlation coefficient matrix (Table 3) revealed that the radial growth of L. gmelinii followed three variation models corresponding to the three latitude levels (low: lower than 49° N; middle: 49° N to 51° N; high: higher than 52° N).

Table 3. Correlation matrix for the standard ring-width chronologies for the seven sampling sites, 1960–2013 (all values significant at \( p < 0.05 \)).

<table>
<thead>
<tr>
<th></th>
<th>AES</th>
<th>BKT</th>
<th>WEQH</th>
<th>YTLH</th>
<th>MEDG</th>
<th>MG</th>
<th>MH</th>
</tr>
</thead>
<tbody>
<tr>
<td>AES</td>
<td>1</td>
<td>0.46</td>
<td>0.44</td>
<td>0.43</td>
<td>0.26</td>
<td>0.06</td>
<td>0.01</td>
</tr>
<tr>
<td>BKT</td>
<td>1</td>
<td>0.31</td>
<td>0.63</td>
<td>0.36</td>
<td>0.36</td>
<td>0.36</td>
<td>0.23</td>
</tr>
<tr>
<td>WEQH</td>
<td>1</td>
<td>0.52</td>
<td>0.68</td>
<td>0.44</td>
<td>0.26</td>
<td></td>
<td></td>
</tr>
<tr>
<td>YTLH</td>
<td>1</td>
<td>0.39</td>
<td>0.46</td>
<td>0.21</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>MEDG</td>
<td>1</td>
<td>0.39</td>
<td>0.29</td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>MG</td>
<td>1</td>
<td>0.63</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>MH</td>
<td>1</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

4. Discussion

Dendroecological and dendroclimatological methods were used to analyze the climate–growth relationships for Larix gmelinii stands in the western Greater Khingan Mountains of China. In the past few decades, large-spatial-scale studies of tree species have been conducted at mid- to high latitudes of the Northern Hemisphere. For example, in the eastern Canadian boreal forest, Jianguo Huang explored the responses of four major boreal tree species to climatic variations along a broad latitudinal gradient and concluded that the growth of Betula papyrifera stands located north of 49°N and the growth of P. mariana and P. banksiana stands north of 47° N were positively correlated with winter and spring temperatures, whereas the growth of stands at more southerly latitudes was negatively correlated with winter and spring temperatures [8]. In Central and Northern Europe, Mäkinen reported that the effect of low temperatures on tree radial growth was more significant at northern sites, while the relative importance of precipitation was greater for the southern sites [36]. Lloyd investigated the climate response of three Siberian taiga species, Picea obovata, Larix cajanderi, and Pinus sylvestris across a latitudinal gradient in central Siberia and concluded that climate response varied among and within populations of all species, and positive responses to temperature prevailed at northern sites, where trees grew faster in years with warm early summers. Negative responses to warming declined along the south to north latitudinal gradient [37].

In the current study, the redundancy analysis (RDA) and correlation analysis revealed that Larix gmelinii exhibited some consistent radial growth–climate associations across the latitudes (Figure 4, Figure 5). Temperatures during the current June to July proved to be the most important environmental factor affecting the radial growth of Larix gmelinii at the regional scale. Meanwhile, in China, the increase of summer temperatures is most significant in the northeast [19]. Our results indicate that higher monthly maximum temperatures (Tmax) in the current June may increase water stress, limiting the radial growth of Larix gmelinii across the gradient, and that serious growth inhibition occurred at the mid- and low-latitude sites. Our results also suggest that a warmer June monthly minimum temperatures (Tmin) in the previous year favored radial growth in the current growing season across the gradient, especially at AES, YTLH, and MG. The highest precipitation in our study region occurs in summer, and a warmer June Tmin most likely enhances photosynthetic rates and accelerates radial growth [38]. At the same time, more sugars are stored in tree biomass over the winter, which are available for tree growth in the following summer [39]. However, hot summer temperatures might enhance stand respiration and evapotranspiration, thus resulting in increased water deficits [8,40] that limit the growth of Larix gmelinii. In addition, warm July temperatures in the current year can significantly limit the radial growth of Larix gmelinii north of 52° N. The inhibition of radial growth at
higher latitudes by warm July temperatures may be due to the low precipitation and/or shallowness of the soil in this region. In addition, these sites have sandy soils, likely resulting in low soil water-holding capacity, so that the warmer temperature will cause a soil water deficit at these latitudes. This supports the findings of Wang [21], who also documented negative relationships between *Larix gmelinii* radial growth and the current-year July temperatures in this area.

Current-year February–March monthly mean temperatures (Tm) and monthly minimum temperatures (Tmin) were found to negatively affect the radial growth of *Larix gmelinii* across latitudes, particularly at MH and MG (high latitudes) and AES (low latitude). Our results are consistent with those of Wang [19], who also observed that the temperature in the previous November and December and in the March of the current year had a negative effect on the growth of *Larix gmelinii* at MH. Tree-ring radial growth is known to be influenced by the amount of storage compounds and the current soil moisture regime [41]. This is likely due to warmer winter (December–February) temperatures, which increase the loss of carbohydrates stored for growth in the following summer [42]. More detailed investigation on the influence of late-winter and early-spring temperatures on radial growth is needed to gain a deeper understanding of the relationship between radial increment and climatic variation.

The warm Tm in April-May was favorable for *Larix gmelinii* radial growth at all sites, especially favoring radial growth at BKT and WEQH. Warm spring temperatures can result in early leaf emergence or budburst and thus result in better growth due to an extended growing season. Graumlich reported that warm temperatures in April may favor coniferous species because their evergreen foliage allows them to take advantage of the warmth during the early growing season [43]. Similarly, Suni emphasized the significance of air temperature triggering the initiation of photosynthesis in boreal conifers in spring [44]. Goodine further noted the particular importance of spring air temperature for tree species whose root systems grow in frozen soils to trigger the initiation of photosynthesis because for tree species whose root systems extend below the frost line, the onset of photosynthesis may be triggered by soil thawing [38]. More unfrozen soil water could be present because of the warm spring air temperatures that are necessary for the onset of photosynthesis. The photosynthetic rate of well-watered foliage may also be favored by higher air temperatures [38]. Our results are in line with those of Wang [21], who suggested that *Larix gmelinii* radial growth at high latitudes (52°N to 53°N) was positively correlated with air temperature during the current-year April to May. The radial growth of *Larix gmelinii* at the mid-latitude sites (especially MEDG and WEQH) was significantly and positively affected by Tm in the previous autumn. However, at the low- and high-latitude sites, a warmer autumn Tm clearly resulted in a prolonged growing season, which led to enhanced growth of *L. gmelinii*.

Our results suggested that precipitation was also an important factor limiting the radial growth in the Greater Khingan Mountains. At almost all latitudes, the growth of *Larix gmelinii* was limited by water availability, as suggested by the significant positive correlation between radial growth and precipitation during the current March to August and the previous June to October. As water supply during the spring season is crucial for tree growth, sufficient accumulation of snow and the consequent melting may meet the water demands of *Larix gmelinii* stands during an early growing season [45]. Summer precipitation has a direct influence on the tree-ring increment because rain is used to almost immediately recharge the soil water reservoir. Low summer precipitation may cause water stress and reduced photosynthesis [46]. The correlation between ring width and summer precipitation in mid-high Northern Hemisphere sites is predominantly positive [47]. In line with this, the ring width in this study was significantly positively correlated with summer precipitation across most of the latitudes in the Greater Khingan Mountains, but at WEQH (mid-latitude), the association with summer precipitation was negative. A possible reason is that on the steep slopes, considerable amounts of soil nutrients are washed out by the high summer precipitation. Ring-width correlations with precipitation during the growing season at northern MEDG were non-significant and were weaker than those for the southern latitudes. This might be because the relative importance of precipitation during the
growing season increased in the warmer southern region. This study revealed that wet winters have a relatively weak negative impact on ring widths at almost all latitudes. A possible reason for this may be that deep snow during winter could delay snow-melt and limit tree growth [48]. Our results are in line with those of Wang [19], who showed that the radial growth of *Larix gmelinii* from 52° N to 53° N was negatively correlated with winter precipitation because precipitation from the previous winter was important for bud protection. Li reported that the summer and autumn precipitation in Northeastern China has decreased since the 1960s [49]; thus, precipitation’s limiting effect increases for *Larix gmelinii* growth. Based on the findings in this study, rapid warming and drying were not conducive to *L. gmelinii* growth in the Greater Khingan Mountains. This is consistent with the results that have suggested that *L. gmelinii* in the Greater Khingan Mountains would be replaced by other species [16–18].

5. Conclusions

Temperatures during current June to July were found to be the dominant environmental factor affecting the radial growth of *Larix gmelinii* in the Greater Khingan Mountains. More specifically, *Larix gmelinii* growth might be decreasing in radial growth by higher Tmax in June, especially in the stands at low and middle latitudes. *Larix gmelinii* radial growth at high latitudes (MG and MH) might be reduced by warmer temperatures in July. Meanwhile, precipitation was also an important factor limiting the radial growth at the regional scale. *Larix gmelinii* might be decreasing in radial growth from decreasing precipitation occurring across the gradient. Our results show that there is a decreasing trend in *Larix gmelinii* radial growth under the observed general increase of temperature in the Greater Khingan Mountains in recent years.

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References


