



<https://helda.helsinki.fi>

Helda

A pH dependent sulfate formation mechanism caused by hypochlorous acid in the marine atmosphere

Liu, Jiarong

Elsevier B.V.

2021-09-15

Liu, J, Ning, A, Liu, L, Wang, H, Kurten, T & Zhang, X 2021, 'A pH dependent sulfate formation mechanism caused by hypochlorous acid in the marine atmosphere', *The Science of the Total Environment*, vol. 787, 147551. <https://doi.org/10.1016/j.scitotenv.2021.147551>

<http://hdl.handle.net/10138/357778>

[10.1016/j.scitotenv.2021.147551](https://doi.org/10.1016/j.scitotenv.2021.147551)

cc_by_nc_nd

acceptedVersion

Downloaded from Helda, University of Helsinki institutional repository.

This is an electronic reprint of the original article.

This reprint may differ from the original in pagination and typographic detail.

Please cite the original version.

1 **A pH dependent sulfate formation mechanism caused by**
2 **hypochlorous acid in the marine atmosphere**

3 Jiarong Liu¹, An Ning¹, Ling Liu¹, Huixian Wang², Theo Kurtén^{3*}, Xiuhui Zhang^{1*}

4 ¹ Key Laboratory of Cluster Science, Ministry of Education of China, School of Chemistry and Chemical
5 Engineering, Beijing Institute of Technology, Beijing 100081, China.

6 ² Beijing Guodian Longyuan Environment Engineering Co. Ltd, Beijing 100081, China.

7 ³ Department of Chemistry, University of Helsinki, Helsinki FI-00014, Finland

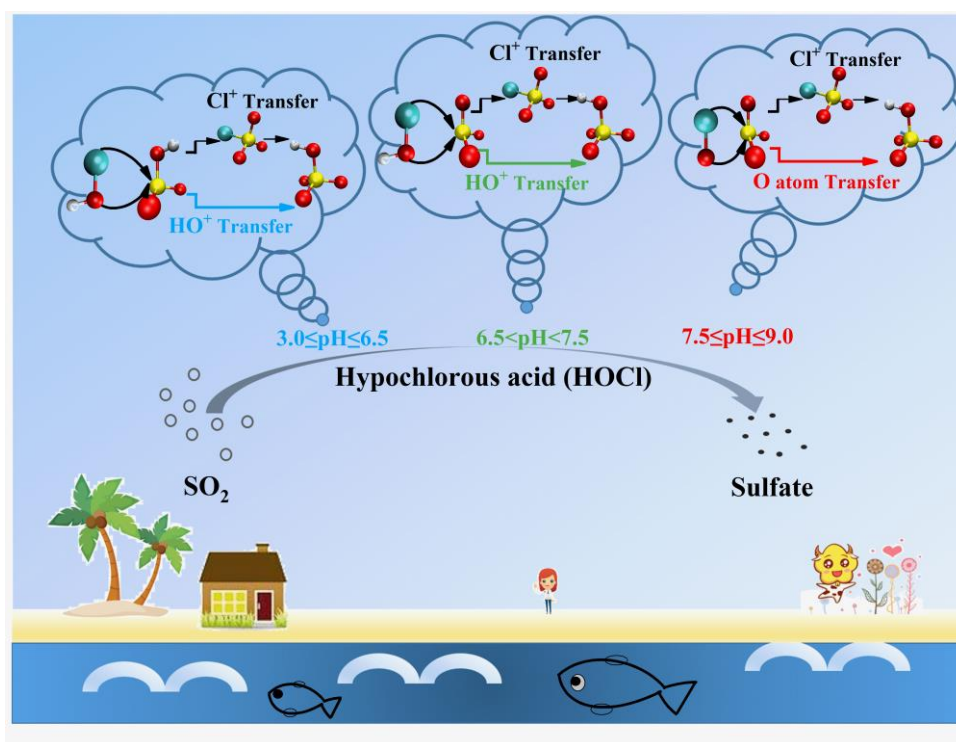
8 * Corresponding authors, E-mail addresses: theo.kurten@helsinki.fi, zhangxiuhui@bit.edu.cn.

9 **Abstract**

10 Secondary sulfate plays a crucial role in forming marine aerosol, which in turn is an
11 important source of natural aerosol at a global level. Recent experimental studies
12 suggest that oxidation of S(IV) compounds, in practice dissolved sulfur dioxide, to
13 sulfate (S(VI)) by hypochloric acid could be one of the most significant pathways for
14 sulfate formation in marine areas. However, the exact mechanism responsible for this
15 process remains unknown. Using high-level quantum chemical calculations, we studied
16 the reaction between dissolved sulfur dioxide and hypochloric acid. We account for the
17 dominant protonation states of reactants in the pH range 3.0-9.0. We also consider
18 possible catalytic effects of species such as H₂O. Our results show that sulfate formation
19 in HOCl+HOSO₂⁻ and HOCl+SO₃²⁻ reactions relevant to acidic and nearly neutral
20 conditions can occur either through previously proposed Cl⁺ transfer or through a novel
21 HO⁺ transfer mechanism. In alkaline conditions, where the dominant reactants are OCl⁻
22 and SO₃²⁻, an O atom transfer mechanism proposed in previous experimental studies

23 may be more important than Cl^+ transfer. Catalysis by common cloud-water species is
24 found to lower barriers of Cl^+ transfer mechanisms substantially. Nevertheless, we find
25 that the dominant $\text{S(IV)}+\text{HOCl}$ reaction mechanism for the full studied pH range is
26 HO^+ transfer from HOCl to SO_3^{2-} , which leads directly to sulfate formation without
27 ClSO_3^- intermediates. The rate-limiting barrier of this reaction is low, leading to an
28 essentially diffusion-controlled reaction rate. S(IV) lifetimes due to this reaction
29 decrease with increasing pH due to the increasing fractional population of SO_3^{2-} .
30 Especially in neutral and alkaline conditions, depletion of HOCl by the reaction is so
31 rapid that S(IV) oxidation will be controlled mainly by mass transfer of gas-phase HOCl
32 to the liquid phase. The mechanism proposed here may help to explain marine sulfate
33 sources missing from current atmospheric models.

34 **Keywords:** Marine aerosol formation; Oxidation pathway; Aqueous phase reaction;
35 DFT studies



37 **1. Introduction**

38 It is well known that aerosol particles have significant influences on the
39 environment, global climate, and human health(Almeida et al., 2013; Dumka et al.,
40 2020; Ehn et al., 2014; Wang et al., 2015; Zhao et al., 2020). However, there are still
41 some large uncertainties in the estimation of the effects of aerosol particles on all these
42 issues. A major reason for this uncertainty is our poor understanding of the formation
43 process, and therefore the abundance, of aerosol particles especially in marine
44 areas(Choi et al., 2017; Li et al., 2019). Excluding primary sea spray emissions, marine
45 aerosol mainly originates from the secondary formation processes, and is one of the
46 most important sources of natural aerosols at a global level(Dowd et al., 2004; O'Dowd
47 and de Leeuw, 2007; Yan et al., 2020). A series of measurements have shown that the
48 total mass of marine aerosols is dominated by sulfate, with a mass fraction of between
49 50% and 67%(Choi et al., 2017; Huang et al., 2018). As a predominant component of
50 atmospheric aerosol in general, sulfate is known to have profound impact not only on
51 marine climate, but also on the global climate system(Li et al., 2017; Zhang et al., 2019).
52 On one hand, sulfate aerosol have a direct cooling effect on the radiative balance by
53 efficiently scattering solar radiation(Gen et al., 2019; Huang et al., 2014). On the other
54 hand, sulfate aerosol also act as the potential cloud condensation nuclei (CCN), thus
55 altering the number, size and water uptake of the cloud droplets, and leading to even
56 greater, though highly uncertain, indirect radiative effects(Hung et al., 2018).
57 Measurements have shown that CCN number concentrations dramatically increase
58 following increases of sulfate aerosol(Novakov et al., 1994).

59 The main source of secondary (non-sea salt) sulfate in marine areas is oxidation of
60 sulfur dioxide (SO₂), which mainly originates from the oxidation of dimethyl sulfide
61 (DMS) emitted by oceanic phytoplankton (Chen et al., 2018; Hung and Hoffmann, 2015;
62 Landim et al., 2018) or from long-range transport of human emissions including fossil
63 fuel combustion and industrial processes (Alexander et al., 2009). The oxidation of SO₂
64 to sulfate has attracted much interest, and very many studies have been performed to
65 identify the detailed formation process of sulfate in different conditions (Chen et al.,
66 2016; Gen et al., 2019; He et al., 2014; Li et al., 2018; Wang et al., 2016; Wang et al.,
67 2020). Based on model simulations, the main sulfate (S(VI)) formation process in
68 marine areas is thought to be oxidation of dissolved SO₂ (H₂SO₃ + HOSO₂⁻ + SO₃²⁻) in
69 cloud water. Dissolved SO₂ is often denoted simply as S(IV). This process generally
70 prevails over gas-phase oxidation due to the high relative humidity (RH), and
71 consequent high liquid water content, in typical marine air (Alexander et al., 2012; Chen
72 et al., 2016). The main known oxidation mechanisms of S(IV) in cloud water are its
73 reactions with hydrogen peroxide (H₂O₂), ozone (O₃), nitrogen dioxide (NO₂) and
74 various transition metal ions together with molecular oxygen (O₂) (Alexander et al.,
75 2009; Alexander et al., 2012; Brandt and van Eldik, 1995; Chandler et al., 1988; Clifton
76 et al., 1988; Harris et al., 2013; Hung and Hoffmann, 2015; Kunen et al., 1983; McArdle
77 and Hoffmann, 1983; Shen et al., 2012). All the corresponding oxidation mechanisms
78 are pH dependent, for example due to the pH dependence of the relative concentrations
79 of the three reactant species H₂SO₃, HOSO₂⁻ and SO₃²⁻. Although all the above sulfate
80 formation processes have been considered in atmospheric models, marine sulfate

81 concentrations are still underestimated significantly. This underestimation is often
82 attributed to the existence of some important but as yet unknown pathways for
83 converting S(IV) to sulfate(Cheng et al., 2016; Huang et al., 2014; Wang et al., 2014;
84 Wang et al., 2016; Zheng et al., 2015).

85 Recent studies have suggested that S(IV) oxidation reaction caused by hypochloric
86 acid(HOCl) in cloud water is a potentially important source of sulfate, as hypochloric
87 acid is present in the marine atmosphere (with reported gas-phase mixing ratios ranging
88 from 5 to 173 pptv), and is well-known to be an efficient oxidant(Alexander et al., 2009;
89 Alexander et al., 2012; Chen et al., 2016; Chen et al., 2017; Keene et al., 1998; Vogt et
90 al., 1996; von Glasow et al., 2002). In particular, Chen *et al.* provided observational
91 evidence for the role of hypochloric acid in sulfate formation in marine areas by a
92 combination of observations and modeling of the oxygen isotope ratio
93 $\Delta^{17}\text{O}(\text{nssSO}_4^{2-})$ (Chen et al., 2016). They found that the oxidation processes of S(IV)
94 by hypohalous acids (HOX =HOCl, HOBr) can contribute from 33 to 50 % of the total
95 sulfate production over the entire area sampled. Similarly, Keene *et al.* determined that
96 reaction with hypochloric acid is the principal oxidation pathway for S(IV) leading to
97 sulfate formation at a pH value of 5.5 using the photochemical box model MOCCA in
98 marine areas(Keene et al., 1998).

99 Laboratory experiments have also been performed to investigate the dynamics and
100 mechanisms of the corresponding reaction processes. Liu *et al.* directly determined the
101 rate of sulfate formation from the reaction between HOCl and HOSO_2^- in low-pH
102 aerosol particles by measuring the concentration of sulfate as a function of time in a

103 flow tube. They proposed a possible two-step sulfate formation mechanism involving
104 Cl^+ transfer(Liu and Abbatt, 2020). Fogelman *et al.* obtained the reaction rate constant
105 between HOCl and SO_3^{2-} at near-neutral pH by measuring the loss of reactant per unit
106 time, and also proposed a Cl^+ transfer mechanism(Fogelman et al., 1989). Despite these
107 studies, the speculated pathway for sulfate formation has not been completely identified
108 or verified, because the key chlorosulfate ion (ClSO_3^-) intermediate of the Cl^+ transfer
109 mechanism has not been directly characterized in the experiments. Despite the
110 electrostatic repulsion inevitably present between two anions, Fogelman *et al.* found
111 that the $\text{OCl}^- + \text{SO}_3^{2-}$ reaction can lead to sulfate formation in basic conditions. They
112 proposed that the reaction proceeds via a one-step mechanism involving O atom
113 transfer(Fogelman et al., 1989). Even though the speculated O atom transfer mechanism
114 seems reasonable, the corresponding product has also not been experimentally
115 identified. A Cl^+ transfer mechanism could be possible also in basic conditions, as this
116 mechanism would involve a O^{2-} leaving group, which might have an even greater
117 potential to react spontaneously with H_2O (and thus promote the reaction) than the Cl^-
118 leaving group of the Cl^+ transfer mechanism.

119 As both dissolved S(IV) and hypochloric acid exist in different protonation states
120 depending on the pH (see Fig. 1 below for details), and the different forms of the
121 reactants are likely to display distinct reaction kinetics, the overall S(IV) + HOCl
122 reaction mechanism is likely to be complex. Furthermore, other molecules, including
123 both water and the product sulfate, may participate in the reaction mechanisms, again
124 depending on the pH. In order to elucidate the overall reaction for S(IV) + HOCl, we

125 have accordingly used high-level quantum chemical methods to explore several
126 possible sulfate-forming processes for the wide pH range of 3.0-9.0(Fridlind and
127 Jacobson, 2000; Pszenny et al., 2003) relevant to marine areas. We have also probed
128 the catalytic effects of common atmospheric species such as water and sulfate on the
129 S(IV) + HOCl reactions.

130 **2. Methods**

131 **Electronic structure calculation.** Density functional theory (DFT) calculations were
132 carried out using the Gaussian 16, revision A. 03 programs package(Frisch et al., 2016).
133 The geometry optimizations and vibrational frequency calculations of reactants,
134 products, and transition states were performed at the M06-2X/6-311++G (3df, 3pd)
135 level of theory(Elm et al., 2012; Mardirossian and Head-Gordon, 2016; Zhao and
136 Truhlar, 2008), which has been found to be well suited to compute thermochemistry
137 and barrier heights of molecular systems relevant to the atmosphere(Li et al., 2018;
138 Rong et al., 2020; Zhang et al., 2018; Zhao and Truhlar, 2008). The connection of each
139 transition state to reactants and products was confirmed by intrinsic reaction coordinate
140 (IRC) calculations at the same level of theory. Single point energy corrections were
141 calculated at the CCSD(T)-F12/cc-pVDZ-F12 level of theory with the ORCA 4.10
142 program, using the geometry optimized at the M06-2X/6-311++G (3df, 3pd) level of
143 theory(Lane and Kjaergaard, 2009; Neese, 2012; Peterson et al., 2008). In addition, the
144 solvent environment was modelled using the continuum solvation model based on
145 solute electron density (SMD) at the M06-2X/6-311++G (3df, 3pd) level of
146 theory(Marenich et al., 2009). SMD is considered to be highly adequate to describe

147 detailed reaction processes and predict the corresponding reaction free energy barriers
148 in aqueous solutions(Miguel et al., 2016; Ostovari et al., 2018; Xu and Coote, 2019),
149 especially for atmospherically relevant processes(Keshavarz et al., 2021; Minakata et
150 al., 2015). Also, some studies suggest that the SMD-based implicit solvent treatment
151 may even outperform explicit water treatment in predicting free energy barriers of
152 reactions in aqueous environment(Chen et al., 2019). Free energies reported here
153 include solvent corrections, and correspond to a temperature of 298.15K, and a
154 reference concentration of 1 mol/L. The details of the free energy calculations are
155 summarized in Section 1.1 of the SI.

156 **Kinetics calculations.** The studied mechanisms involve both (barrierless) complex
157 or intermediate formation steps, and subsequent reactions of these complexes or
158 intermediates. Our kinetics treatment is based on the pseudo-steady-state approach
159 (PSSA), and contains two types of parameters: usually bimolecular equilibrium
160 constants (K_{eq}), and formally unimolecular reaction rate coefficients (k). The latter are
161 calculated in terms of Transition State Theory (TST)(Evans and Polanyi, 1935; Eyring,
162 1935; Truhlar et al., 1983) ignoring tunneling as the reactions involve heavy-atom
163 transfers. These parameters are calculated as follows:

$$164 \quad K_{\text{eq}}=(c_0)^{\Delta n} \exp \left(-\frac{\Delta G_{\text{eq}}}{RT}\right) \quad (1)$$

$$165 \quad k=\frac{k_{\text{B}} T}{h} \exp \left(-\frac{\Delta G^{\ddagger}}{RT}\right) \quad (2)$$

166 where ΔG_{eq} is the difference of the Gibbs free energy between the reactant(s) and the
167 product(s) of a reaction, R is the gas constant, T is the temperature, k_{B} is the Boltzmann
168 constant, h is the Planck constant, c_0 is the reference concentration (1 mol/L) at which

169 ΔG_{eq} is calculated, and Δn is the change in the number of separate species in the
170 reaction (-1 for the typical case of a bimolecular association reaction where two
171 reactants form a complex or an adduct). Similarly, ΔG^0 is the Gibbs free energy barrier,
172 i.e the difference in the Gibbs free energy between a reactant (usually a reactant
173 complex or adduct) and the transition state.

174 The bimolecular rate constant for an overall reaction (k_{bim}) involving the formation
175 of an initial complex, which then needs to cross a barrier, is then:

$$176 \quad k_{\text{bim}} = K_{\text{eq}} \times k \quad (3)$$

177 In addition, diffusion limitations should be considered especially for low-barrier
178 or barrierless reactions. Collins-Kimball theory (Collins and Kimball, 1949) was used
179 to calculate the overall rate constants k_{overall} for reactions involving an initial barrierless
180 complex formation:

$$181 \quad k_{\text{overall}} = \frac{k_{\text{bim}} \times k_{\text{D}}}{k_{\text{bim}} + k_{\text{D}}} \quad (4)$$

182 Where k_{D} is the steady-state Smoluchowski rate constant that can be expressed as Eq.
183 (5) (Chang, 2000; Smoluchowski, 1917)

$$184 \quad k_{\text{D}} = 4\pi R D_{\text{AB}} N_{\text{A}} \quad (5)$$

185 where N_{A} is the Avogadro constant and R denotes the reaction distance. Generally, R is
186 the sum of the radii of two reactants for free diffusion processes. In addition, D_{AB} is the
187 mutual diffusion coefficient of the reactants A and B, computed as the sum of D_{A} and
188 D_{B} (Truhlar, 1985), which in turn are calculated by using the Stokes-Einstein
189 approach (Einstein, 1905; Stokes, 1903):

$$190 \quad D_{\text{A or B}} = \frac{k_{\text{B}} T}{6\pi\eta a_{\text{A or B}}} \quad (6)$$

191 where η is the viscosity of water, and $a_{A \text{ or } B}$ is the radius of reactant A or B, which can
192 be estimated, assuming spherical reactants, using the volume of the reactants computed
193 with Gaussian 16, revision A. 03 (Frisch et al., 2016). Rates for the barrierless
194 dissociation of complexes were computed by detailed balance as k_D/K_{eq} , with k_D and
195 K_{eq} computed from Eq (5) and Eq (1), respectively.

196 In addition, it should be noted that, in the present system, reaction might occur
197 between two anions, and thus, the diffusion in a coulomb potential has an R value
198 smaller than that of free diffusion process because of the strong electrostatic repulsion.
199 k_D might then decrease because of the decreased R . In this case, R can be calculated
200 as (Roussel, 1985):

$$201 \quad R = \frac{Z_A Z_B e^2}{4\pi\epsilon k_B [e^{\frac{Z_A Z_B e^2}{4\pi\epsilon k_B r}} - 1]} \quad (7)$$

202 where Z_A and Z_B are the charges for the two ions, ϵ is the dielectric constant for water,
203 k_B is Boltzmann's constant, T is the temperature, and r is the sum of radii of the two
204 reactants.

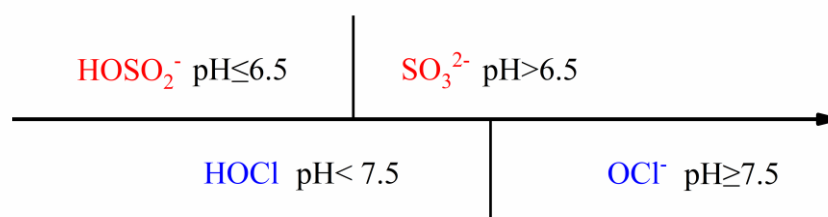
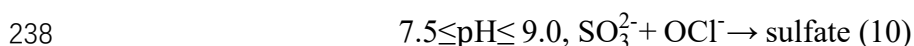
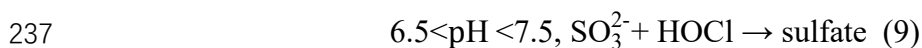
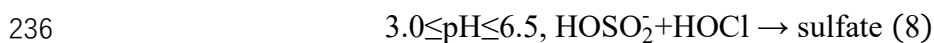
205 **Wavefunction analysis.** Wavefunction analysis was carried out at the start of the study
206 to determine the possible reaction mechanisms available to each combination of
207 reactants. The electrostatic potential (ESP) on the molecular van der Waals (vdw)
208 surface is crucial for predicting the potential reaction sites of intermolecular interactions.
209 The site possessing more positive ESP tends to attract the site which possesses more
210 negative ESP, and vice versa (Manzetti and Lu, 2013; Murray and Politzer, 2011).
211 We accordingly calculated the ESP color-mapped isosurface using Multiwfn 3.7 (Lu
212 and Chen, 2012), and VMD 1.9.3 (Humphrey et al., 1996). Furthermore, the contour

213 map of electron density difference between transition states and reactant complexes,
214 which can indicate the structural changes in the reaction process, was analyzed using
215 Multiwfn 3.7(Lu and Chen, 2012).

216 **3. Results and discussion**

217 The reactants S(IV) and hypochloric acid exist in different protonation states at
218 different pH, potentially leading to different reaction mechanisms. As shown in Fig. 1,
219 based on measured pK_a values, the S(IV) reactant exists primarily in the form of
220 HOSO_2^- at low (but atmospherically relevant) pH, and in the form of SO_3^{2-} in near-
221 neutral or basic conditions ($\text{pH} > 6.5$)(C. Drexler et al., 1992; Kunen et al., 1983). H_2SO_3
222 is the main form of S(IV) only at a very low pH ($\text{pH} < 1.5$)(C. Drexler et al., 1992). Also,
223 this form is extremely unstable(Li and McKee, 1997; Voegelé et al., 2002; Voegelé et
224 al., 2004). Thus, considering the range of pH from 3.0 to 9.0 in the marine areas, and
225 the instability of H_2SO_3 , the corresponding oxidation reactions involving H_2SO_3 were
226 not considered in the present work. In addition, it should be noted that although two
227 bisulfite ion isomers (HSO_3^- and HOSO_2^-) have been found in the aqueous phase, only
228 the dominant one with the proton bonded to the oxygen atom has been chosen for the
229 study of oxidation reactions in the present work(Baird and Taylor, 1981; Brown and
230 Barber, 1995; Horner and Connick, 1986). For the hypochloric acid reactant, HOCl is
231 the dominant form at $\text{pH} < 7.5$, while OCl^- dominates for $\text{pH} > 7.5$ (Shaka et al., 2007).
232 See Section 1.2 of the SI for details. To comprehensively investigate the reaction
233 process of S(IV) oxidation by hypochloric acid, we have considered multiple reactions

234 mechanisms as suggested by ESP surfaces, involving all the dominant reactant forms
235 for a wide range of pH, as illustrated by Eqs. (8), (9) and (10):



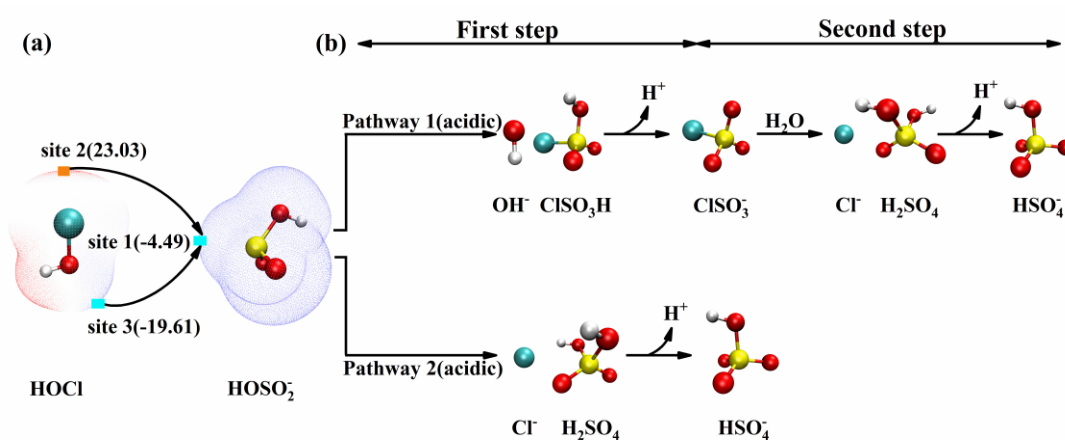
239

240 **Fig. 1.** The dominant forms of reactants S(IV) and hypochlorous acid at different ranges of pH in the
241 cloud water.

242 **3.1. Sulfate formation mechanism in weakly acidic water ($3.0 \leq \text{pH} \leq 6.5$)**

243 Based on the above analysis, S(IV) and hypochlorous acid will be mainly exist as
244 HOSO_2^- and HOCl , respectively, at the typical pH ($\text{pH}=3.0-6.5$) of cloud water. In order
245 to identify the potential reaction sites between HOSO_2^- and HOCl , an in-depth
246 exploration of the ESP on the molecular vdw surface has been performed, and shown
247 in Fig. 2(a). The site which possesses more positive ESP has stronger ability to attract
248 electrophiles, and has a greater ability to react with the site possessing more negative
249 ESP(Lu and Chen, 2012; Manzetti and Lu, 2013; Murray and Politzer, 2011). As
250 illustrated in Fig. 2(a), the chlorine atom (site 2) of HOCl possesses more positive ESP,
251 implying that it has the potential to interact with the sulfur atom (site 1) of HOSO_2^- with
252 negative ESP. Similarly, the oxygen atom (site 3) of HOCl could also interact with the

253 sulfur atom (site 1) of HOSO_2^- , because of the much more negative ESP of site 3
 254 compared to site 1. Thus, based on the ESP surface, there are two potential reaction
 255 pathways of HOSO_2^- with HOCl to form sulfate. In addition to the previously proposed
 256 Cl^+ transfer mechanism (Pathway 1(acidic))(Liu and Abbatt, 2020), the ESP results
 257 suggest that also a hitherto unreported HO^+ transfer mechanism (Pathway 2(acidic))
 258 should be investigated.



259

260 **Fig. 2.** (a) ESP-mapped molecular vdW surfaces of HOCl and HOSO_2^- , in units of kcal/mol.
 261 Surface local minima and maxima of ESPs of different functional groups in the corresponding
 262 molecules are represented as blue and orange squares, respectively. (b) Two possible pathways for
 263 the reaction of HOSO_2^- with HOCl leading to sulfate formation. The white, cyan, red, and yellow
 264 spheres represent H, Cl, O, and S atoms, respectively.

265 The potential energy profiles for the reaction pathways discovered for $\text{HOSO}_2^- +$
 266 HOCl are shown in Figs. 3(a) and (b). These figures also show the optimized structures
 267 of the reactant complexes (R_{cn}), transition states (TS_n) and product complexes (P_{cn}) for
 268 the two studied pathways. As shown in Fig. 3 (a), the reaction of HOSO_2^- with HOCl
 269 via Pathway 1(acidic) begins with the formation of reactant complex $\text{HOCl}\dots\text{HOSO}_2^-$

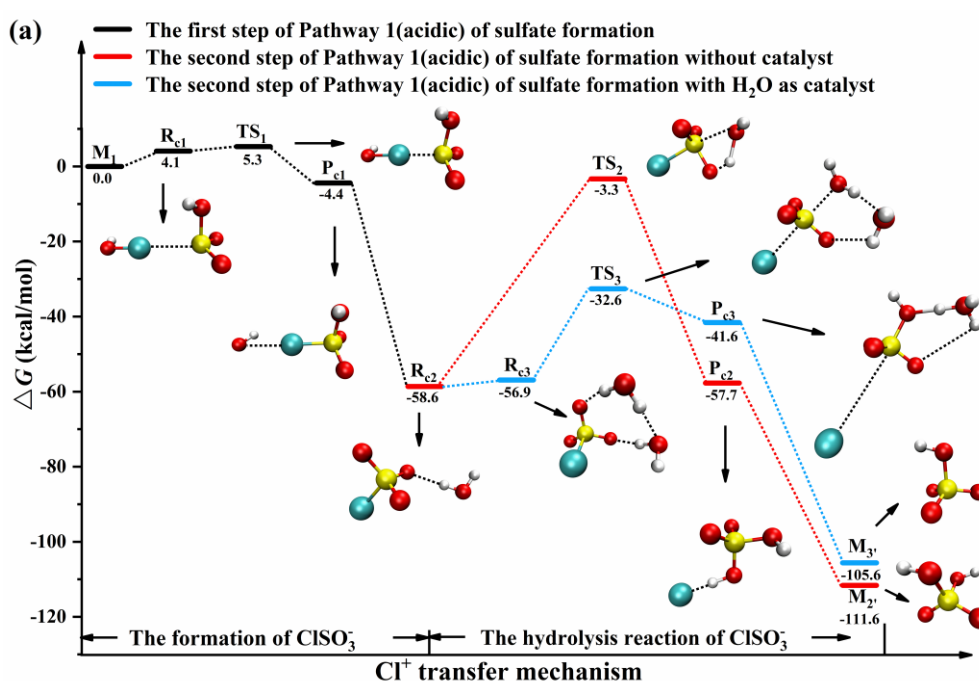
270 (R_{c1}). This complex then crosses a shallow barrier (5.3 kcal/mol relative to the separated
271 reactants) corresponding to TS_1 , forming a complex of $ClSO_3H...OH^-$, (denoted P_{c1}).
272 Next, proton transfer forms a $ClSO_3^-...H_2O$ complex (denoted R_{c2}), lowering the free
273 energy by 54.2 kcal/mol. (This complex is bound by 2.8 kcal/mol with respect to
274 separated $ClSO_3^- + H_2O$; not shown in the Fig. 3(a)) The next step on reaction Pathway
275 1(acidic) is hydrolysis of the $ClSO_3^-$ intermediate. As shown in Fig. 3(a), the
276 $ClSO_3^-...H_2O$ complex (R_{c2}) is connected to a four-membered ring transition state TS_2
277 (55.3 kcal/mol higher in free energy), where one proton of H_2O transfers to the oxygen
278 atom of $ClSO_3^-$, while the remaining hydroxyl of H_2O binds to the sulfur atom of $ClSO_3^-$.
279 The product of this reaction step is a complex between sulfuric acid and chloride,
280 $H_2SO_4...Cl^-$ (P_{c2}), which is bound by 53.9 kcal/mol with respect to separated H_2SO_4
281 and Cl^- . Due to the very high free energy of TS_2 , this process is not a competitive route
282 for sulfate formation.

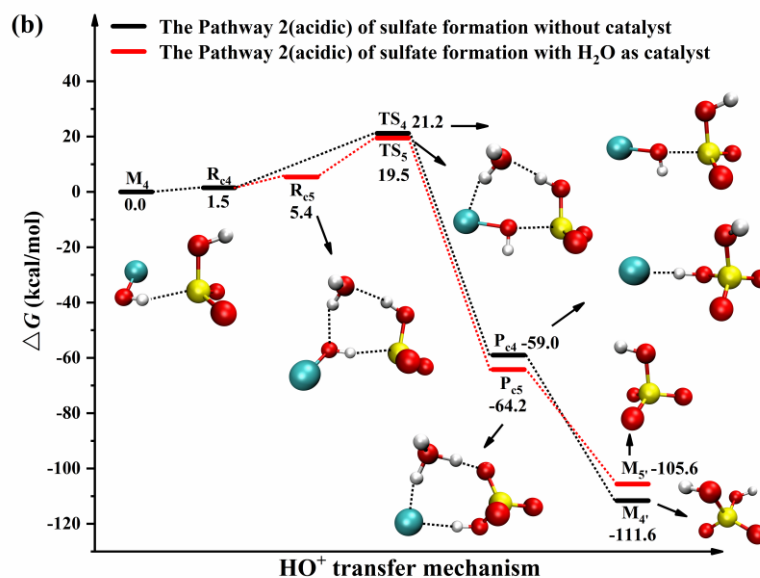
283 However, multiplied hydrogen-bonding species present in cloud water, including
284 water (H_2O), sulfuric acid (H_2SO_4) and bisulfate ions (HSO_4^-), have the ability to
285 catalyze proton transfer reactions(Liu et al., 2015; Meijer and Sprik, 1998; Torrent-
286 Sucarrat et al., 2012). We have investigated the effect of these three species on the
287 hydrolysis of $ClSO_3^-$, and found that all three substantially lower the free energy barrier
288 for the process. The free energy barriers of the catalyzed reactions, relative to the
289 corresponding three-body reactant complexes $ClSO_3^-...H_2O...X$, are 24.3, 25.0, and 23.9
290 kcal/mol, for $X= H_2O, H_2SO_4,$ and HSO_4^- , respectively. See also Fig. 3(a) and Fig. S1
291 of the SI. Here, we discuss further the hydrolysis $ClSO_3^-$ catalyzed by H_2O , the most

292 abundant species in the aqueous environment, as an example of this general catalytic
293 mechanism. As shown in Fig. 3(a), the hydrolysis reaction catalyzed by H₂O proceeds
294 through the clustering of an additional water molecule with ClSO₃⁻...H₂O reactant
295 complex (R_{c2}), forming ClSO₃⁻...H₂O...H₂O (R_{c3}). Next, R_{c3} can be converted into the
296 product complex HSO₄⁻...H₃O⁺...Cl⁻ (denoted P_{c3}), and finally to separated HSO₄⁻,
297 H₃O⁺, and Cl⁻, via a six-membered ring transition state TS₃. The barrier for this
298 catalyzed process is almost 30 kcal/mol lower than that of the uncatalyzed reaction. The
299 reason for this lowering is likely that the strong ring tension of the closed four-
300 membered ring in the uncatalyzed reaction is reduced in the catalyzed reaction, as the
301 ring size is increased by a hydrogen bonding network involving H₂O. The contour maps
302 of the electron density differences between R_{c3} and TS₃ in Fig. S3 illustrate the detailed
303 reaction process. The proton transfer from the reactant H₂O to the oxygen atom of the
304 catalyst H₂O forms a new H-O single bond. The remaining hydroxyl moiety of the
305 reactant H₂O binds to the sulfur atom of ClSO₃⁻, forming a new S-O single bond.
306 Meanwhile, the chloride ion leaves the sulfur atom of ClSO₃⁻ through the cleaving of
307 the Cl-S single bond. Sulfate formation from HOSO₂⁻ + HOCl through the Cl⁺ transfer
308 mechanism (Pathway 1(acidic)) proposed in experimental studies(Liu and Abbatt, 2020)
309 is thus feasible, but only with the help of a catalyst such as H₂O.

310 Moreover, we also discovered an unexpected novel mechanism for HOSO₂⁻ +
311 HOCl, illustrated by Pathway 2(acidic) in Fig. 3(b). Starting with the reactant complex
312 HOSO₂⁻...HOCl (R_{c4}), the reaction goes on to form the product complex H₂SO₄...Cl⁻
313 (P_{c4}) through the transition state TS₄. This mechanism involves the direct formation of

314 sulfate by transferring a HO^+ group from HOCl to HOSO_2^- . The free energy barrier of
 315 this pathway is 21.2 kcal/mol relative to the separated reactants. We investigated
 316 whether also this pathway may be catalyzed by additional hydrogen-bonding molecules.
 317 We find that, H_2O , the most abundant molecule in aqueous environment, can form
 318 hydrogen-bonded complexes with HOSO_2^- , and also with $\text{HOSO}_2^- \dots \text{HOCl}$ (R_{c4}),
 319 leading to a $\text{H}_2\text{O} \dots \text{HOSO}_2^- \dots \text{HOCl}$ reactant complex (R_{c5}). As illustrated in Fig. 3(b),
 320 the catalyzed reaction then leads to the formation of a $\text{H}_3\text{O}^+ \dots \text{HSO}_4^- \dots \text{Cl}^-$ product
 321 complex (P_{c5}) through transition state TS_5 , in which the Cl-O bond of HOCl is cleaved.
 322 and an O-S bond between HOSO_2^- and HOCl is formed. The Gibbs free energy barrier
 323 of this process is reduced to 19.5 kcal/mol relative to the separated reactants, about 1.5
 324 kcal/mol lower than the uncatalyzed reaction. Thus, H_2O may have a modest catalytic
 325 effect also on the HO^+ transfer mechanism (pathway 2(acidic)) discovered here. In
 326 addition, we found that products of this pathway, such as H_2SO_4 and HSO_4^- , could also
 327 act as catalysts. The corresponding processes are shown in Fig. S2 of the SI.





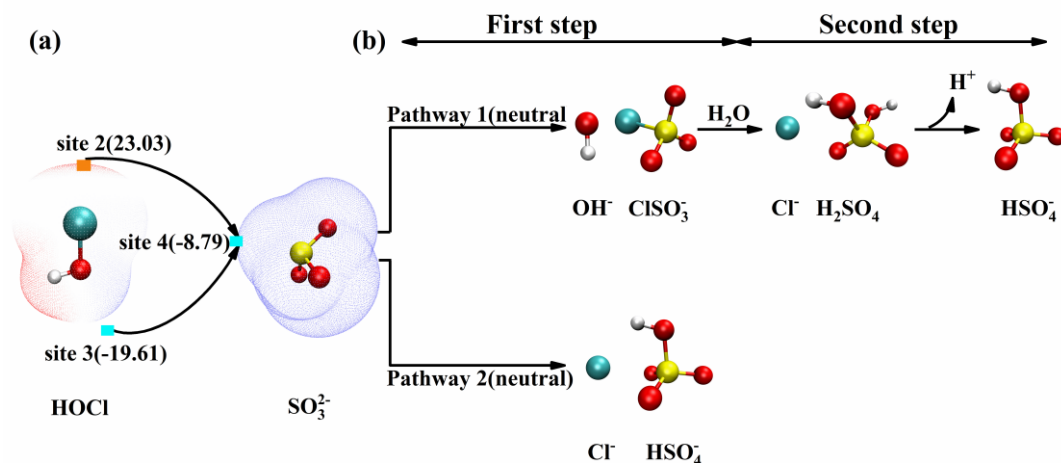
329

330 **Fig. 3.** The Gibbs free energy profile (kcal/mol, at 298.15 K and a reference concentration of 1
 331 mol/L) for the reaction of HOCl with HOSO₂⁻ through (a) Pathway 1(acidic) and (b) Pathway
 332 2(acidic), at the CCSD(T)-F12/cc-pVDZ-F12//M06-2X/6-311++G (3df, 3pd) level of theory.
 333 Solvent corrections are included in the M06-2X calculations using the SMD model. The white, cyan,
 334 red, and yellow spheres represent H, Cl, O, and S atoms, respectively.

335 3.2. Sulfate formation mechanism in neutral water (6.5<pH<7.5)

336 As pH increases, SO₃²⁻ becomes to the dominant form of S(IV). In neutral cloud
 337 water (6.5<pH<7.5), the S(IV) + HOCl reaction thus mainly proceeds through SO₃²⁻ +
 338 HOCl. In addition to the Cl⁺ transfer mechanism (Pathway 1(neutral)) proposed by
 339 earlier experimental studies(Liu and Abbatt, 2020), which is similar to the mechanism
 340 in weakly acidic water except for the participation of SO₃²⁻ instead of HOSO₂⁻, our ESP
 341 analysis (see Fig. 4) suggests the existence of a HO⁺ transfer mechanism also for SO₃²⁻
 342 + HOCl (Pathway 2(neutral)). According to the ESP analysis, the chlorine atom (site 2)
 343 of HOCl possesses positive ESP, enabling it to react with the sulfur atom (site 4) of

344 SO_3^{2-} with negative ESP. Meanwhile, the oxygen atom (site 3) of HOCl may also react
 345 with the sulfur atom (site 4) of SO_3^{2-} because of the much more negative ESP of site 3
 346 compared to site 4.



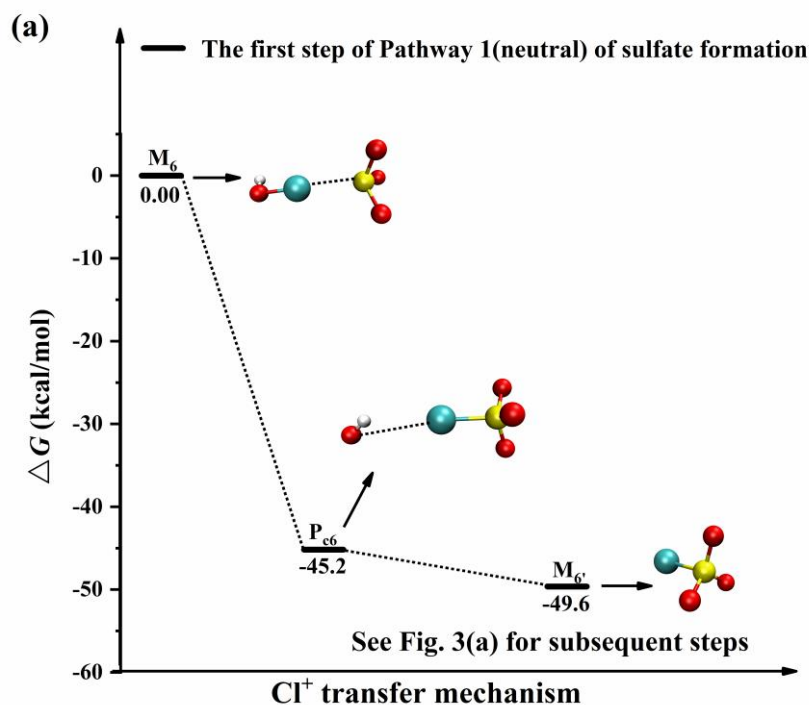
347

348 **Fig. 4.** (a) ESP-mapped molecular vdW surface of HOCl and SO_3^{2-} , in units of kcal/mol. Surface
 349 local minima and maxima of ESPs of different functional groups in the corresponding molecules
 350 are represented as blue and orange squares, respectively. (b) Two possible reaction pathways for the
 351 reaction of SO_3^{2-} with HOCl leading to sulfate formation. The white, cyan, red, and yellow spheres
 352 represent H, Cl, O, and S atoms, respectively.

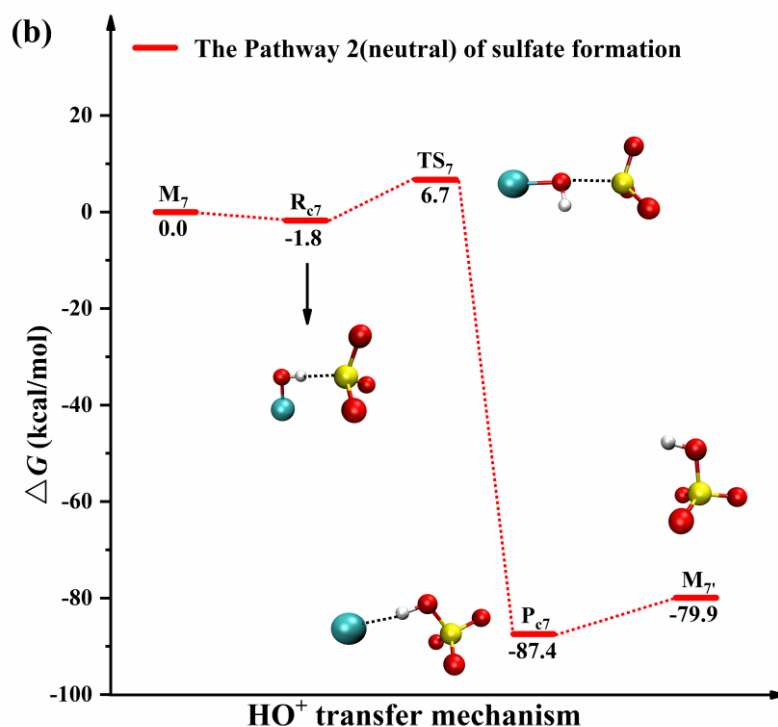
353 The potential energy profiles of Pathway 1(neutral) and Pathway 2(neutral) as well
 354 as the corresponding geometries on the reaction path are shown in Figs. 5(a) and (b),
 355 respectively. In Pathway 1(neutral), similar to the $\text{HOSO}_2^- + \text{HOCl}$ reaction, the first
 356 step involves the barrierless transfer of Cl^+ from HOCl to the S atom of SO_3^{2-} , forming
 357 the ClSO_3^- intermediate, plus a OH^- ion which likely leaves the reacting system due to
 358 electrostatic repulsion (M_6^-). The lack of a barrier suggests that the rate of this reaction
 359 step is controlled by diffusion in the cloud water (see Eq (5)), and is therefore likely to
 360 be very rapid. Our calculations thus verify the efficient formation of the ClSO_3^-

361 intermediate which has not yet been characterized experimentally. ClSO_3^- will
362 subsequently react with H_2O to form sulfate through the hydrolysis reaction already
363 presented in Fig. 3(a), with catalytic effects e.g. from another H_2O molecule as
364 discussed above.

365 The HO^+ transfer mechanism (Pathway 2(neutral) in Fig. 5(b)) begins with the
366 formation of a $\text{HOCl}\dots\text{SO}_3^{2-}$ reactant complex (labelled R_{c7}), which then crosses a fairly
367 low barrier (6.7 kcal/mol relative to the isolated reactant; labelled TS_7), leading directly
368 to HSO_4^- and Cl^- . This low-barrier pathway, which was not considered in previous
369 investigations, is likely the most competitive mechanism for sulfate formation in near-
370 neutral conditions. (As discussed below, it is actually the dominant mechanism for all
371 conditions.)



372



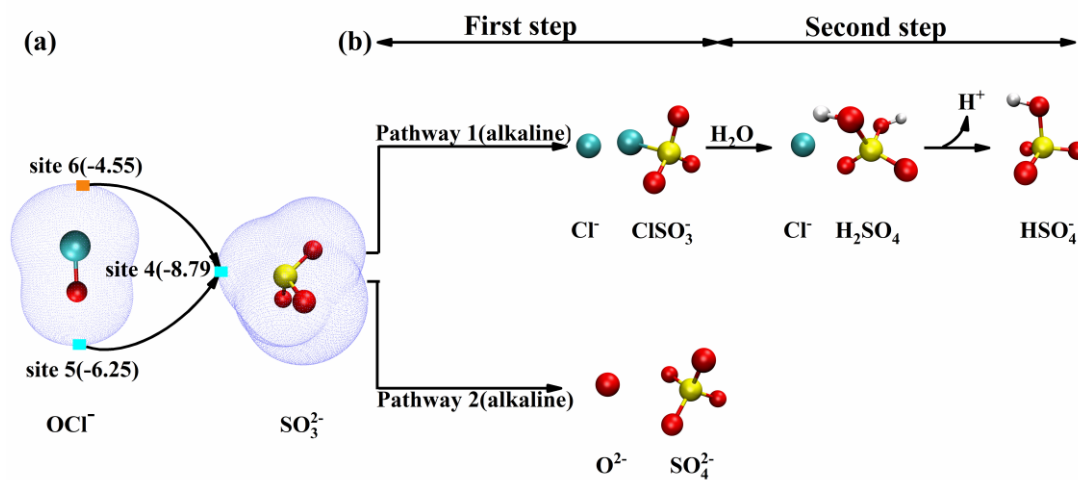
373

374 **Fig. 5.** The activation Gibbs free energy profile (kcal/mol, at 298.15 K and a reference concentration
 375 of 1 mol/L) for the reaction of HOCl with SO₃²⁻ through the (a) Pathway 1(neutral) and the (b)
 376 Pathway 2(neutral) at the CCSD(T)-F12/cc-pVDZ-F12//M06-2X/6-311++G (3df, 3pd) level of
 377 theory. Solvent corrections are included in the M06-2X calculations using the SMD model. The
 378 white, cyan, red, and yellow spheres represent H, Cl, O, and S atoms, respectively.

379 **3.3. Sulfate formation mechanism in alkaline water (7.5 ≤ pH ≤ 9.0)**

380 In the marine areas, when the cloud water is mixed with fresh sea-salt spray, the
 381 atmospheric liquid phase may occasionally be alkaline rather than acidic. With the
 382 increase of pH, the fraction of HOCl decreases, and OCl⁻ becomes to the dominant form
 383 of hypochloric acid. Therefore, in the alkaline cloud water, the oxidation of S(IV) by
 384 hypochloric acid could take place via a OCl⁻ + SO₃²⁻ reaction. While hindered somewhat
 385 by electrostatic repulsion (accounted for here using a modified version of the diffusion

386 rate coefficient, see Eq (7)), this anion-anion reaction has already been proposed to be
 387 feasible in experimental studies(Fogelman et al., 1989). In addition to the O atom
 388 transfer mechanism (Pathway 2(alkaline)) proposed in experimental studies(Fogelman
 389 et al., 1989), ESP results indicate that a Cl^+ transfer mechanism (Pathway 1(alkaline))
 390 analogous to that found in weakly acidic or neutral conditions may also be possible in
 391 alkaline conditions. This mechanism is driven by the attraction between the chloride
 392 atom of OCl^- and the sulfur atom of SO_3^{2-} , as illustrated in the ESP analysis shown in
 393 Fig. 6. Here, the chlorine atom (site 6) of OCl^- possesses more positive ESP, giving it
 394 potential to react with the sulfur atom (site 4) of SO_3^{2-} with more negative ESP. We
 395 have thus investigated both O and Cl^+ transfer mechanisms also for alkaline conditions.



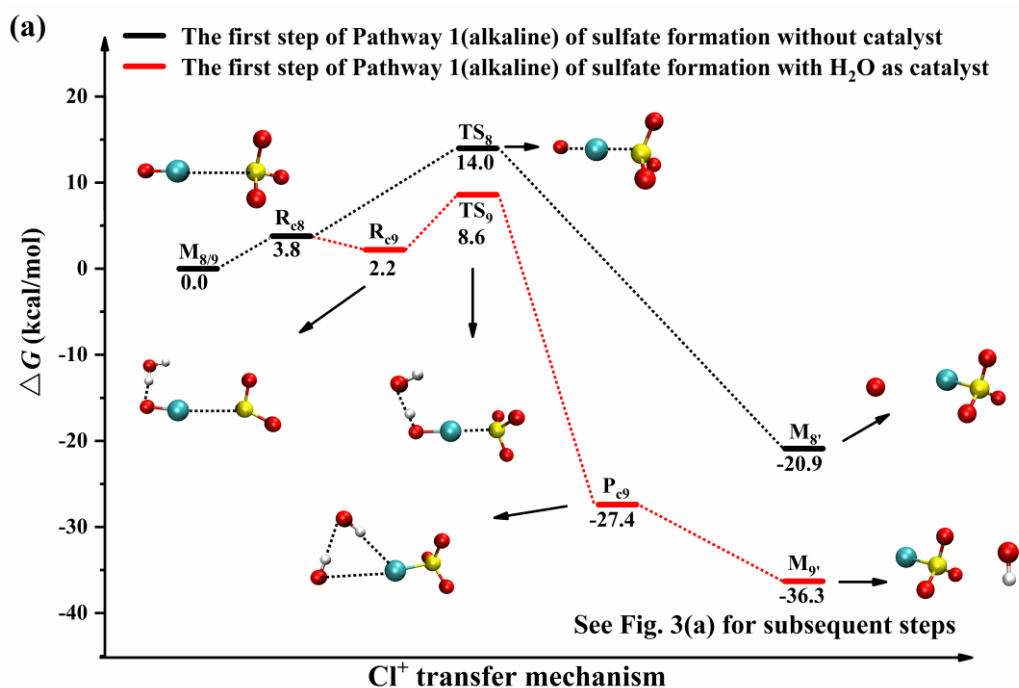
396

397 **Fig. 6.** (a) ESP-mapped molecular vdW surface of OCl^- and SO_3^{2-} , in units of kcal/mol. Surface
 398 local minima and maxima of ESPs of different functional groups in the corresponding molecules
 399 are represented as blue and orange squares, respectively. (b) Two possible reaction pathways for the
 400 reaction of SO_3^{2-} with OCl^- leading to sulfate formation. The cyan, red, and yellow spheres represent
 401 Cl, O, and S atoms, respectively.

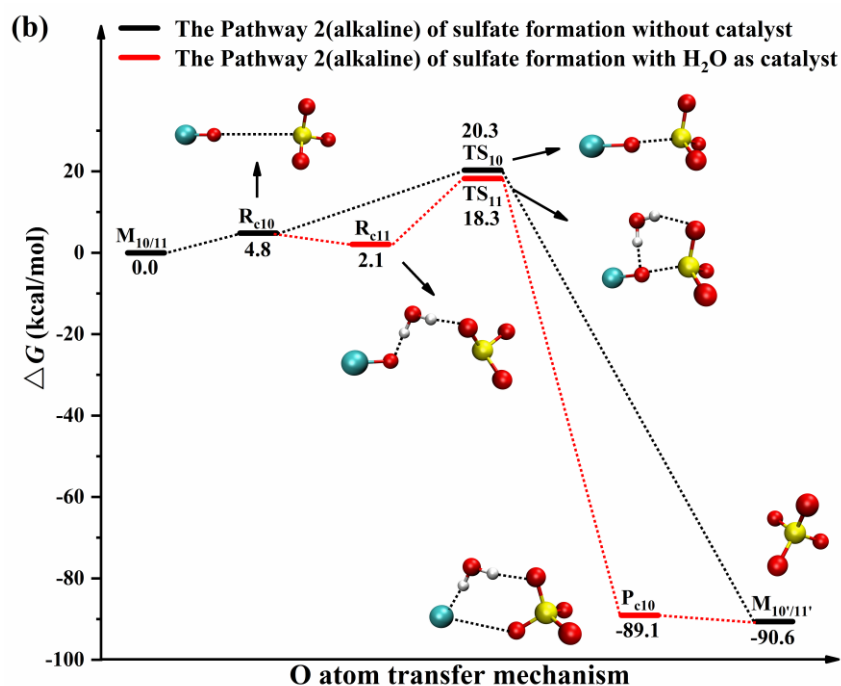
402 The potential energy profiles and the corresponding geometries on the reaction
403 path from OCl^- and SO_3^{2-} are shown in Figs. 7(a) and (b), respectively. Pathway
404 1(alkaline) begins with the formation of a $\text{SO}_3^{2-}\dots\text{OCl}^-$ complex (R_{c8}). Despite the
405 electrostatic repulsion between two anions, our solvent-corrected calculations predict
406 that this complex is only 3.8 kcal/mol above the separated reactants in free energy. R_{c8}
407 is connected to a linear transition state TS_8 , leading to the formation of the ClSO_3^-
408 intermediate and O^{2-} (M_8) with a free energy barrier of 14.0 kcal/mol. ClSO_3^- then
409 reacts with H_2O to form sulfate as shown in Fig. 3(a). Analogous to the acidic and
410 neutral cases, Pathway 1(alkaline) can also be catalyzed by a H_2O molecule. The first
411 step on this pathway is the formation of $\text{SO}_3^{2-}\dots\text{OCl}^-\dots\text{H}_2\text{O}$ complex (R_{c9}), which is
412 connected to transition state TS_9 . In TS_9 , the Cl^+ of OCl^- interacts with the S atom of
413 SO_3^{2-} , forming the ClSO_3^- intermediate (which again hydrolyses as shown in Fig. 3(a)).
414 Meanwhile, the O^{2-} leaving group of OCl^- accepts a proton from H_2O , producing two
415 hydroxyls (OH^-) ions, much more stable products than the O^{2-} formed in the
416 uncatalyzed mechanism. The free energy barrier for the catalyzed mechanism is
417 accordingly 5.4 kcal/mol lower than that of the uncatalyzed mechanism. Although the
418 formation of ClSO_3^- intermediates in the Cl^+ transfer mechanism in alkaline conditions
419 was not considered in previous studies(Fogelman et al., 1989), our calculations indicate
420 that it is highly like to form on both the catalyzed and uncatalyzed reaction pathways.

421 Pathway 2(alkaline) of the $\text{OCl}^- + \text{SO}_3^{2-}$ reaction, shown in Fig. 7(b), also starts
422 with the formation of a $\text{SO}_3^{2-}\dots\text{OCl}^-$ complex (R_{c10}). R_{c10} differs from R_{c9} in the relative
423 orientation of OCl^- with respect to SO_3^{2-} , as shown in Fig. 7(b). R_{c10} is connected to

424 transition state TS₁₀, with a free energy barrier of 20.3 kcal/mol relative to the isolated
 425 reactants, leading directly to the SO₄²⁻ and Cl⁻ product monomers (labelled M₁₀). This
 426 pathway can also be catalyzed by H₂O. The catalyzed reaction starts with the formation
 427 of a SO₃²⁻...OCl⁻...H₂O reactant complex (R_{c11}), which is converted into a
 428 SO₄²⁻...Cl⁻...H₂O product complex (P_{c11}) via the transition state TS₁₁. The free energy
 429 barrier of this process (relative to the isolated reactants) is 2.0 kcal/mol lower than that
 430 of the uncatalyzed pathway. Analogous to Pathway 2(acidic), we thus predict a modest
 431 catalytic effect of water on Pathway 2(alkaline). The contour maps of the electron
 432 density difference between reactant complex R_{c11} and transition state TS₁₁ (see Fig. S3)
 433 indicate that the oxygen atom of OCl⁻ transfers to the sulfur atom of SO₃²⁻, forming a
 434 new O-S bond, while the remaining chloride ion is cleaved from the oxygen atom of
 435 OCl⁻.



436



437

438 **Fig. 7.** The Gibbs free energy profile (kcal/mol, at 298.15 K and a reference concentration of 1

439 mol/L) for the reaction of OCl⁻ with SO₃²⁻ through (a) Pathway 1(alkaline) and (b) Pathway

440 2(alkaline), at the CCSD(T)-F12/cc-pVDZ-F12//M06-2X/6-311++G (3df, 3pd) level of theory.

441 Solvent corrections are included in the M06-2X calculations using the SMD model The white, cyan,

442 red, and yellow spheres represent H, Cl, O, and S atoms, respectively.

443 **4. Atmospheric Implication**

444 Tables 1, 2 and 3 show the S(IV) oxidation rates, and corresponding lifetimes,

445 estimated for the six different pathways, applying the pseudo-steady-state

446 approximation to all intermediates, and using a constant aqueous-phase HOCl

447 concentration of 3.43×10^{-9} M. This concentration corresponds to an aqueous phase in

448 Henry's law equilibrium at 298.15 K with a gas phase containing 0.005 ppbv HOCl.

449 While this gas-phase mixing ratio represents the lower end of the experimentally

450 reported range (0.005 ppbv - 0.173ppbv), our assumption of a Henry's law equilibrium

451 correspondingly represents an upper-limit estimate, which likely does not hold
 452 especially at higher pH as discussed below. We also note that while the pseudo-
 453 unimolecular S(IV) oxidation rates and S(IV) lifetimes by definition do not depend on
 454 the total S(IV) concentration, the rates and lifetimes of each individual channel still
 455 depend on the fractional populations of the different S(IV) reactants (HOSO₂⁻ and
 456 SO₃²⁻). As these are pH – dependent, also the rates and lifetimes vary with the pH. The
 457 alkaline pathways have a further pH – dependence through the concentration of the
 458 OCl⁻ reactant. For details of the rate and lifetime calculations, see the Supplementary
 459 Information.

460 Table 1. Calculated S(IV) oxidation rates ($v_{1\text{acid, cat-S(IV)}}$ and $v_{2\text{acid, cat-S(IV)}}$) and S(IV)
 461 lifetimes for Pathway 1(acidic) and Pathway 2(acidic) from HOSO₂⁻ and HOCl with
 462 H₂O as a catalyst, respectively, assuming [HOCl] = 3.43×10^{-9} M (corresponding to a
 463 Henry's law equilibrium with a gas phase HOCl mixing ratio of 0.005ppbv), and [H₂O]
 464 = 55.56 M (corresponding to the known bulk concentration of H₂O).

pH	Pathway 1(acidic)		Pathway 2(acidic)	
	$v_{1\text{acid, cat-S(IV)}}$	Lifetime	$v_{2\text{acid, cat-S(IV)}}$	Lifetime
	(s ⁻¹)	(s)	(s ⁻¹)	(s)
3	5.25×10^{-17}	1.90×10^{16}	9.77×10^{-10}	1.02×10^9
4	5.55×10^{-17}	1.80×10^{16}	1.03×10^{-9}	9.68×10^8
5	5.53×10^{-17}	1.81×10^{16}	1.03×10^{-9}	9.72×10^8
6	5.05×10^{-17}	1.98×10^{16}	9.40×10^{-10}	1.06×10^9
6.5	4.18×10^{-17}	2.39×10^{16}	7.77×10^{-10}	1.29×10^9

7	2.70×10^{-17}	3.70×10^{16}	5.02×10^{-10}	1.99×10^9
7.5	1.27×10^{-17}	7.85×10^{16}	2.37×10^{-10}	4.22×10^9
8	4.77×10^{-18}	2.10×10^{17}	8.88×10^{-11}	1.13×10^{10}
9	5.17×10^{-19}	1.93×10^{18}	9.62×10^{-12}	1.04×10^{11}

465 Table 2. The calculated S(IV) oxidation rates ($v_{1\text{neu, cat-S(IV)}}$ and $v_{2\text{neu-S(IV)}}$) and the S(IV)
466 lifetimes for Pathway 1(neutral) (with H₂O as a catalyst) and Pathway 2(neutral) from
467 SO₃²⁻ and HOCl, respectively, assuming [HOCl] = 3.43×10^{-9} M (corresponding to a
468 Henry's law equilibrium with a gas phase HOCl mixing ratio of 0.005ppbv), and [H₂O]
469 =55.56 M (corresponding to the known bulk concentration of H₂O).

pH	Pathway 1(neutral)		Pathway 2(neutral)	
	$v_{1\text{neu, cat-S(IV)}}$	Lifetime	$v_{2\text{neu-S(IV)}}$	Lifetime
	(s ⁻¹)	(s)	(s ⁻¹)	(s)
3	5.79×10^{-20}	1.73×10^{19}	2.59×10^{-5}	3.87×10^4
4	6.12×10^{-19}	1.63×10^{18}	2.73×10^{-4}	3.66×10^3
5	6.09×10^{-18}	1.64×10^{17}	2.72×10^{-3}	3.67×10^2
6	5.57×10^{-17}	1.80×10^{16}	2.49×10^{-2}	4.02×10^1
6.5	1.46×10^{-16}	6.87×10^{15}	6.51×10^{-2}	1.54×10^1
7	2.98×10^{-16}	3.36×10^{15}	1.33×10^{-1}	7.52
7.5	4.44×10^{-16}	2.25×10^{15}	1.98×10^{-1}	5.04
8	5.26×10^{-16}	1.90×10^{15}	2.35×10^{-1}	4.25
9	5.70×10^{-16}	1.75×10^{15}	2.55×10^{-1}	3.93

471 Table 3. The calculated S(IV) oxidation rates ($v_{1\text{alk, cat-S(IV)}}$ and $v_{2\text{alk, cat-S(IV)}}$) and S(IV)
 472 lifetimes for Pathway 1(alkaline) and Pathway 2(alkaline) from SO_3^{2-} and OCl^- with
 473 H_2O as a catalyst, respectively, assuming $[\text{HOCl}] = 3.43 \times 10^{-9}$ M (corresponding to a
 474 Henry's law equilibrium with a gas phase HOCl mixing ratio of 0.005ppbv), and $[\text{H}_2\text{O}]$
 475 $= 55.56$ M (corresponding to the known bulk concentration of H_2O).

pH	Pathway 1(neutral)		Pathway 2(neutral)	
	$v_{1\text{alk, cat-S(IV)}}$	Lifetime	$v_{2\text{alk, cat-S(IV)}}$	Lifetime
	(s^{-1})	(s)	(s^{-1})	(s)
3	1.71×10^{-26}	5.85×10^{25}	1.43×10^{-16}	7.00×10^{15}
4	1.81×10^{-24}	5.53×10^{23}	1.51×10^{-14}	6.63×10^{13}
5	1.80×10^{-22}	5.56×10^{21}	1.50×10^{-12}	6.65×10^{11}
6	1.64×10^{-20}	6.08×10^{19}	1.37×10^{-10}	7.28×10^9
6.5	1.36×10^{-19}	7.35×10^{18}	1.14×10^{-9}	8.80×10^8
7	8.79×10^{-19}	1.14×10^{18}	7.34×10^{-9}	1.36×10^8
7.5	4.15×10^{-18}	2.41×10^{17}	3.47×10^{-8}	2.89×10^7
8	1.55×10^{-17}	6.44×10^{16}	1.30×10^{-7}	7.70×10^6
9	1.68×10^{-16}	5.94×10^{15}	1.41×10^{-6}	7.11×10^5

476 Our results indicate that the HO^+ transfer mechanism (Pathway 2(neutral)) from
 477 $\text{SO}_3^{2-} + \text{HOCl}$ completely dominates sulfate formation through the $\text{HOCl} + \text{S(IV)}$
 478 reaction at all atmospherically relevant pH, despite the relatively low fractional
 479 populations of SO_3^{2-} at low pH and HOCl at high pH. Even at a pH of 9, the second-
 480 fastest reaction mechanism (Pathway 2(alkaline)) is still over five orders of magnitude

481 slower than Pathway 2(neutral).

482 The explanation for this is simple: Pathway 2(neutral) has a much lower rate-
483 limiting barrier than any of the other channels, leading to an essentially diffusion-
484 limited rate constant. Pathways 1(acidic), 1(neutral) and 1(alkaline) all involve a ClSO_3^-
485 hydrolysis step, which has a substantial barrier even after catalysis by a third body such
486 as H_2O (see Fig. 3(a)). Pathways 2(acidic) and 2(alkaline) have lower but still
487 substantial barriers, which are only modestly lowered by catalysts. The rates of both
488 alkaline pathways are further decreased by electrostatic repulsion between the reactant
489 anions.

490 Our simple lifetime estimates show that if HOCl is present in the aqueous phase,
491 the HOCl + S(IV) reaction is likely to be competitive with other known S(IV) oxidation
492 channels for the full studied pH range of 3 to 9. Especially toward the upper end of this
493 range, the computed lifetimes are unrealistically short, as experimentally measured
494 lifetimes for liquid-phase S(IV) are seldom below about 1000 s (Berglen, 2004; Fang et
495 al., 2019; Kasting et al., 1987; Lelieveld and Heintzenberg, 1992; Wang et al., 2020;
496 Zhang et al., 2020). The explanation for this discrepancy is that especially at higher pH,
497 depletion of HOCl through the HO^+ transfer from $\text{SO}_3^{2-} + \text{HOCl}$ reaction (Pathway
498 2(neutral)) is so rapid that the Henry's law equilibrium between the gas phase and the
499 aqueous phase will never have time to form. S(IV) oxidation will then be controlled by
500 the rate of diffusion and dissolution of gas-phase HOCl to the liquid phase. We also
501 note that the rapid depletion of HOCl through the direct HO^+ transfer reaction may also
502 help explain the absence of experimental observations of ClSO_3^- or other intermediates,

503 despite the fact that their formation is also predicted to be rapid.

504 The extremely rapid S(IV) oxidation by HOCl predicted by our calculations may
505 help explain observations of unexpectedly efficient sulfate formation in marine areas.
506 While S(IV) + HOCl reactions have been proposed previously (Fogelman et al., 1989;
507 Liu and Abbatt, 2020), our results reveal that the dominant reaction mechanism is
508 different from previous speculations. Chemistry-transport modelling incorporating the
509 pH-dependent reaction mechanism proposed here, and accounting for the rate-limiting
510 gas-liquid mass transfer of HOCl, will likely be useful for improving regional and
511 global sulfate budgets. Finally, hypobromic acid (HOBr) is chemically very similar to
512 hypochloric acid, and might conceivably also contribute to S(IV) oxidation and thus
513 sulfate formation in marine areas. The S(IV) + HOBr reaction should therefore be
514 explored in the future.

515 **5. Conclusion**

516 We have performed high-level quantum chemical calculations to comprehensively
517 examine the mechanism for sulfate formation through the oxidation of dissolved
518 sulfurous acid (denoted S(IV)) by hypochloric acid over wide range of pH (from 3.0 to
519 9.0). Our results demonstrate that while previously proposed Cl⁺ transfer mechanisms
520 exist, and might be catalyzed by e.g. H₂O molecules, sulfate formation from the S(IV)
521 + HOCl reaction is completely dominated by a novel low-barrier HO⁺ transfer
522 mechanism from HOCl + SO₃²⁻. This reaction mechanism forms sulfate directly,
523 without involving intermediates such as ClSO₃⁻, which likely explains the puzzling lack
524 of observations of such intermediates in experimental studies. The rate of S(IV)

525 oxidation via this reaction increases with pH due to the increasing fractional population
526 of SO_3^{2-} . Even in acidic conditions (pH = 3), the S(IV) lifetime predicted for this
527 reaction, assuming a liquid phase in Henry's law equilibrium with a gas-phase
528 containing 5 pptv HOCl, is less than nine hours. At higher pH, the rate of HOCl
529 depletion is likely so rapid that the Henry's law equilibrium will not hold, and S(IV)
530 oxidation will be controlled by mass transfer of HOCl to the liquid phase. The HOCl +
531 SO_3^{2-} reaction mechanism proposed here is likely to help explain sulfate formation
532 missing from current atmospheric chemistry models, and thus narrow the gap between
533 field observations and modelling on the distribution of sulfate, as well as improve our
534 overall understanding of marine secondary aerosol.

535 **Declaration of competing interest**

536 The authors declare that they have no known competing financial interests or
537 personal relationships that could have appeared to influence the work reported in this
538 paper.

539 **Acknowledgements**

540 We thank the National Natural Science Foundation of China (21976015). L.
541 L. thanks the China Postdoctoral Science Foundation (Grant 2020M680013). T. K.
542 thanks the Academy of Finland for Funding and the CSC – IT Center for Science for
543 computer resources.

544 **CRedit authorship contribution statement**

545 Jiarong Liu: Data curation, Formal analysis, Investigation, Writing - original draft;
546 Writing - review & editing. An Ning and Ling Liu: Supervision; Validation,

547 Visualization. Huixian Wang: Supervision; Validation. Theo Kurtén: Writing - review
548 & editing, Methodology, Formal analysis. Xiuhui Zhang: Conceptualization, Writing -
549 review & editing, Formal analysis, Data curation.

550

551 **6. References:**

- 552 Alexander, B., Park, R. J., Jacob, D. J., Gong, S., 2009. Transition metal-catalyzed oxidation of atmospheric sulfur:
553 Global implications for the sulfur budget. *J. Geophys. Res.*, 114, D2309, <http://doi.org/10.1029/2008JD010486>
- 554 Alexander, B., Allman, D. J., Amos, H. M., Fairlie, T. D., Dachs, J., Hegg, D. A., Sletten, R. S., 2012. Isotopic
555 constraints on the formation pathways of sulfate aerosol in the marine boundary layer of the subtropical
556 northeast Atlantic Ocean. *J. Geophys. Res.*, 117, D6304, <http://doi.org/10.1029/2011JD016773>
- 557 Almeida, J., Schobesberger, S., Kurten, A., Ortega, I. K., Kupiainen-Maatta, O., Praplan, A. P., Adamov, A.,
558 Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Donahue, N. M., Downard, A., Dunne,
559 E., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M.,
560 Henschel, H., Jokinen, T., Junninen, H., Kajos, M., Kangasluoma, J., Keskinen, H., Kupc, A., Kurten, T.,
561 Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Leiminger, M., Leppa, J., Loukonen, V., Makhmutov, V., Mathot,
562 S., McGrath, M. J., Nieminen, T., Olenius, T., Onnela, A., Petaja, T., Riccobono, F., Riipinen, I., Rissanen, M.,
563 Rondo, L., Ruuskanen, T., Santos, F. D., Sarnela, N., Schallhart, S., Schnitzhofer, R., Seinfeld, J. H., Simon,
564 M., Sipila, M., Stozhkov, Y., Stratmann, F., Tome, A., Trostl, J., Tsagkogeorgas, G., Vaattovaara, P., Viisanen,
565 Y., Virtanen, A., Vrtala, A., Wagner, P. E., Weingartner, E., Wex, H., Williamson, C., Wimmer, D., Ye, P.,
566 Yli-Juuti, T., Carslaw, K. S., Kulmala, M., Curtius, J., Baltensperger, U., Worsnop, D. R., Vehkamäki, H.,
567 Kirkby, J., 2013. Molecular understanding of sulphuric acid-amine particle nucleation in the atmosphere.
568 *Nature*, 502, 359-363, <http://doi.org/10.1038/nature12663>
- 569 Baird, N. C., Taylor, K. F., 1981. Ab initio MO calculations for the oxides, oxyacids, and oxyanions of S(IV) and
570 S(VI). *J. Comput. Chem.*, 2, 225-230, <http://doi.org/10.1002/jcc.540020303>
- 571 Berglen, T. F., 2004. A global model of the coupled sulfur/oxidant chemistry in the troposphere: The sulfur cycle. *J.*
572 *Geophys. Res.*, 109, D19310, <http://doi.org/10.1029/2003JD003948>
- 573 Brandt, C., van Eldik, R., 1995. Transition metal-catalyzed oxidation of sulfur (IV) oxides. Atmospheric-relevant
574 processes and mechanisms. *Chem. Rev.*, 95, 119-190, <http://doi.org/10.1021/cr00033a006>
- 575 Brown, R. E., Barber, F., 1995. Ab initio studies of the thermochemistry of the bisulfite and the sulfonate ions and

576 related compounds. *J. Phys. Chem.*, 99, 8071-8075, <http://doi.org/10.1021/j100020a034>

577 C. Drexler, H. E., Fecher, B., Wannowius, K. J., 1992. Kinetics and mechanism of sulfur (IV) oxidation by hydrogen
578 peroxide in aqueous phase: the non-linear parts of the pH-profile. *Bur. Bunsenges. Phys.*, 96, 481-485,
579 <http://doi.org/10.1002/bbpc.19920960349>

580 Chandler, A. S., Choularton, T. W., Dollardt, G. J., Eggleton, A. E. J., Gay, M. J., Hill, T. A., Jones, B. M. R., Tyler,
581 B. J., Bandy, B. J., Penkett, S. A., 1988. Measurements of H₂O₂ and SO₂ in clouds and estimates of their
582 reaction rate. *Nature*, 336, 562-565, <http://doi.org/10.1038/336562a0>

583 Chang, R., 2000. *Physical Chemistry for the Chemical and Biological Sciences*. Sausalito, CA.

584 Chen, J., Shao, Y., Ho, J., 2019. Are explicit solvent models more accurate than implicit solvent models? a case
585 study on the menshutkin reaction. *J. Phys. Chem. A*, 123, 5580-5589, <http://doi.org/10.1021/acs.jpca.9b03995>

586 Chen, Q., Sherwen, T., Evans, M., Alexander, B., 2018. DMS oxidation and sulfur aerosol formation in the marine
587 troposphere: a focus on reactive halogen and multiphase chemistry. *Atmos. Chem. Phys.*, 18, 13617-13637,
588 <http://doi.org/10.5194/acp-2018-410>

589 Chen, Q., Geng, L., Schmidt, J. A., Xie, Z., Kang, H., Dachs, J., Cole-Dai, J., Schauer, A. J., Camp, M. G., Alexander,
590 B., 2016. Isotopic constraints on the role of hypohalous acids in sulfate aerosol formation in the remote marine
591 boundary layer. *Atmos. Chem. Phys.*, 16, 11433-11450, <http://doi.org/10.5194/acp-16-11433-2016>

592 Chen, Q., Schmidt, J. A., Shah, V., Jaeglé, L., Sherwen, T., Alexander, B., 2017. Sulfate production by reactive
593 bromine: Implications for the global sulfur and reactive bromine budgets. *Geophys. Res. Lett.*, 44, 7069-7078,
594 <http://doi.org/10.1002/2017GL073812>

595 Cheng, Y., Zheng, G., Wei, C., Mu, Q., Zheng, B., Wang, Z., Gao, M., Zhang, Q., He, K., Carmichael, G., Poschl,
596 U., Su, H., 2016. Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in
597 China. *Sci Adv*, 2, e1601530, <http://doi.org/10.1126/sciadv.1601530>

598 Choi, Y., Rhee, T. S., Collett, J. L., Park, T., Park, S., Seo, B., Park, G., Park, K., Lee, T., 2017. Aerosol
599 concentrations and composition in the North Pacific marine boundary layer. *Atmos. Environ.*, 171, 165-172,
600 <http://doi.org/10.1016/j.atmosenv.2017.09.047>

601 Clifton, C. L., Altstein, N., Huie, R. E., 1988. Rate constant for the reaction of nitrogen dioxide with sulfur (IV) over
602 the pH range 5.3-13. *Environ. Sci. Technol.*, 22, 586-589, <http://doi.org/10.1021/es00170a018>

603 Collins, F. C., Kimball, G. E., 1949. Diffusion-controlled reaction rates. *J. Colloid Interf. Sci.*, 4, 425-437,
604 [http://doi.org/10.1016/0095-8522\(49\)90023-9](http://doi.org/10.1016/0095-8522(49)90023-9)

605 Dowd, C. D. O., Facchini, M. C., Cavalli, F., Ceburnis, D., Mircea, M., Decesari, S., Fuzzi, S., Yoon, Y. J., Putaud,

606 J., 2004. Biogenically driven organic contribution to marine aerosol. *Nature*, 431, 676-680,
607 <http://doi.org/10.1038/nature02959>

608 Dumka, U. C., Ningombam, S. S., Kaskaoutis, D. G., Madhavan, B. L., Song, H. J., Angchuk, D., Jorphail, S., 2020.
609 Long-term (2008–2018) aerosol properties and radiative effect at high-altitude sites over western trans-
610 Himalayas. *Sci. Total Environ.*, 734, 139354, <http://doi.org/10.1016/j.scitotenv.2020.139354>

611 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R.,
612 Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma,
613 J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G., Canagaratna, M.,
614 Maso, M. D., Berndt, T., Petäjä, T., Wahner, A., Kerminen, V., Kulmala, M., Worsnop, D. R., Wildt, J., Mentel,
615 T. F., 2014. A large source of low-volatility secondary organic aerosol. *Nature*, 506, 476-479,
616 <http://doi.org/10.1038/nature13032>

617 Einstein, A., 1905. On the movement of small particles suspended in stationary liquids required by the molecular-
618 kinetic theory of heat. *Ann. Phys.*, 17, 549-560,

619 Elm, J., Bilde, M., Mikkelsen, K. V., 2012. Assessment of density functional theory in predicting structures and free
620 energies of reaction of atmospheric prenucleation clusters. *J. Chem. Theory Comput.*, 8, 2071-2077,
621 <http://doi.org/10.1021/ct300192p>

622 Evans, M. G., Polanyi, M., 1935. Some applications of the transition state method to the calculation of reaction
623 velocities, especially in solution. *Trans. Faraday Soc.*, 31, 875-893, <http://doi.org/10.1039/TF9353100875>

624 Eyring, H., 1935. The activated complex in chemical reactions. *J. Chem. Phys.*, 3, 107-115,
625 <http://doi.org/10.1063/1.1749604>

626 Fang, Y., Ye, C., Wang, J., Wu, Y., Hu, M., Lin, W., Xu, F., Zhu, T., 2019. Relative humidity and O₃ concentration
627 as two prerequisites for sulfate formation. *Atmos. Chem. Phys.*, 19, 12295-12307, [http://doi.org/10.5194/acp-](http://doi.org/10.5194/acp-19-12295-2019)
628 19-12295-2019

629 Fogelman, K. D., Walker, D. M., Margerum, D. W., 1989. Non-metal redox kinetics: hypochlorite and hypochlorous
630 acid reactions with sulfite. *Inorg. Chem.*, 28, 986-993, <http://doi.org/10.1002/chin.198925018>

631 Fridlind, A. M., Jacobson, M. Z., 2000. A study of gas-aerosol equilibrium and aerosol pH in the remote marine
632 boundary layer during the First Aerosol Characterization Experiment (ACE 1). *J. Geophys. Res. Atmos.*, 105,
633 17325-17340, <http://doi.org/10.1029/2000JD900209>

634 Frisch, M. J., Trucks, G. W., Schlegel, H. B., Scuseria, G. E., Robb, M. A., Cheeseman, J. R., Scalmani, G., Barone,
635 V., Petersson, G. A., Nakatsuji, H., Li, X., Caricato, M., Marenich, A. V., Bloino, J., Janesko, B. G., Gomperts,

636 R., Mennucci, B., Hratchian, H. P., Ortiz, J. V., Izmaylov, A. F., Sonnenberg, J. L., Williams-Young, D., Ding,
637 F., Lipparini, F., Egidi, F., Goings, J., Peng, B., Petrone, A., Henderson, T., Ranasinghe, D., Zakrzewski, V.
638 G., Gao, J., Rega, N., Zheng, G., Liang, W., Hada, M., Ehara, M., Toyota, K., Fukuda, R., Hasegawa, J., Ishida,
639 M., Nakajima, T., Honda, Y., Kitao, O., Nakai, H., Vreven, T., Throssell, K., Montgomery, J. A., Peralta, J. E.,
640 Ogliaro, F., Bearpark, M. J., Heyd, J. J., Brothers, E. N., Kudin, K. N., Staroverov, V. N., Keith, T. A.,
641 Kobayashi, R., Normand, J., Raghavachari, K., Rendell, A. P., Burant, J. C., Iyengar, S. S., Tomasi, J., Cossi,
642 M., Millam, J. M., Klene, M., Adamo, C., Cammi, R., Ochterski, J. W., Martin, R. L., Morokuma, K., Farkas,
643 O., Foresman, J. B., Fox, D. J., 2016. Gaussian, Inc., Wallingford CT. Gaussian, Inc., Wallingford CT,
644 Gen, M., Zhang, R., Huang, D. D., Li, Y., Chan, C. K., 2019. Heterogeneous SO₂ oxidation in sulfate formation by
645 photolysis of particulate nitrate. *Environ. Sci. Technol. Lett.*, 6, 86-91,
646 <http://doi.org/10.1021/acs.estlett.8b00681>
647 Harris, E., Sinha, B., van Pinxteren, D., Tilgner, A., Fomba, K. W., Schneider, J., Roth, A., Gnauk, T., Fahlbusch,
648 B., Mertes, S., Lee, T., Collett, J., Foley, S., Borrmann, S., Hoppe, P., Herrmann, H., 2013. Enhanced role of
649 transition metal ion catalysis during in-cloud oxidation of SO₂. *Science*, 340, 727-730,
650 <http://doi.org/10.1126/science.1230911>
651 He, H., Wang, Y., Ma, Q., Ma, J., Chu, B., Ji, D., Tang, G., Liu, C., Zhang, H., Hao, J., 2014. Mineral dust and NO_x
652 promote the conversion of SO₂ to sulfate in heavy pollution days. *Sci Rep*, 4, 1-5,
653 <http://doi.org/10.1038/srep04172>
654 Horner, D. A., Connick, R. E., 1986. Equilibrium quotient for the isomerization of bisulfite ion from HSO₃⁻ to SO₃H.
655 *Inorg. Chem*, 25, 2414-2417, <http://doi.org/10.1021/ic00234a026>
656 Huang, S., Wu, Z., Poulain, L., van Pinxteren, M., Merkel, M., Assmann, D., Herrmann, H., Wiedensohler, A., 2018.
657 Source apportionment of the submicron organic aerosols over the Atlantic Ocean from 53°N to 53°S using
658 HR-ToF-AMS. *Atmos. Chem. Phys*, 24, 18043-18062, <http://doi.org/10.5194/acp-2018-307>
659 Huang, X., Song, Y., Zhao, C., Li, M., Zhu, T., Zhang, Q., Zhang, X., 2014. Pathways of sulfate enhancement by
660 natural and anthropogenic mineral aerosols in China. *J. Geophys. Res. Atmos.*, 119, 14165-14179,
661 <http://doi.org/10.1002/2014JD022301>
662 Humphrey, W., Dalke, A., Schulten, K., 1996. VMD: Visual molecular dynamics. *J Mol Graph*, 14, 33-38,
663 [http://doi.org/https://doi.org/10.1016/0263-7855\(96\)00018-5](http://doi.org/https://doi.org/10.1016/0263-7855(96)00018-5)
664 Hung, H., Hsu, M., Hoffmann, M. R., 2018. Quantification of SO₂ oxidation on interfacial surfaces of acidic micro-
665 droplets: Implication for ambient sulfate formation. *Environ. Sci. Technol.*, 52, 9079-9086,

666 <http://doi.org/10.1021/acs.est.8b01391>

667 Hung, H., Hoffmann, M. R., 2015. Oxidation of gas-phase SO₂ on the surfaces of acidic microdroplets: implications
668 for sulfate and sulfate radical anion formation in the atmospheric liquid phase. *Environ. Sci. Technol.*, 49,
669 13768-13776, <http://doi.org/10.1021/acs.est.5b01658>

670 Kasting, J. F., Liu, M. S., Toon, O. B., Turco, R. P., 1987. The sulfur cycle in the marine atmosphere. *J. Geophys.*
671 *Res. Atmos.*, 92, 943-963, <http://doi.org/10.1029/JD092iD01p00943>

672 Keene, W. C., Sander, R., Pszenny, A. A. P., Vogt, R., Crutzen, P. J., Galloway, J. N., 1998. Aerosol pH in the
673 marine boundary layer: A review and model evaluation. *J. Aerosol Sci.*, 29, 339-356,
674 [http://doi.org/10.1016/S0021-8502\(97\)10011-8](http://doi.org/10.1016/S0021-8502(97)10011-8)

675 Keshavarz, F., Thornton, J. A., Vehkamäki, H., Kurtén, T., 2021. Reaction mechanisms underlying unfunctionalized
676 alkyl nitrate hydrolysis in aqueous aerosols. *ACS Earth Space Chem.*, 5, 210-225,
677 <http://doi.org/10.1021/acsearthspacechem.0c00253>

678 Kunen, S. M., Lazrus, A. L., Kok, G. L., Heike, B. G., 1983. Aqueous oxidation of SO₂: by hydrogen peroxid. *J*
679 *Geophys Res*, 88, 3671-3674, <http://doi.org/10.1029/JC088iC06p03671>

680 Landim, A. A., Teixeira, E. C., Agudelo-Castañeda, D., Schneider, I., Silva, L. F. O., Wiegand, F., Kumar, P., 2018.
681 Spatio-temporal variations of sulfur dioxide concentrations in industrial and urban area via a new statistical
682 approach. *Air Qual Atmos Health*, 11, 801-813, <http://doi.org/10.1007/s11869-018-0584-2>

683 Lane, J. R., Kjaergaard, H. G., 2009. Explicitly correlated intermolecular distances and interaction energies of
684 hydrogen bonded complexes. *J. Chem. Phys.*, 131, 34307, <http://doi.org/10.1063/1.3159672>

685 Lelieveld, J., Heintzenberg, J., 1992. Sulfate cooling effect on climate through in-cloud oxidation of anthropogenic
686 SO₂. *Science*, 258, 117-120, <http://doi.org/10.1126/science.258.5079.117>

687 Li, G., Bei, N., Cao, J., Huang, R., Wu, J., Feng, T., Wang, Y., Liu, S., Zhang, Q., Tie, X., Molina, L. T., 2017. A
688 possible pathway for rapid growth of sulfate during haze days in China. *Atmos. Chem. Phys.*, 17, 1-43,
689 <http://doi.org/10.5194/acp-2016-994>

690 Li, H., Zhong, J., Vehkamäki, H., Kurtén, T., Wang, W., Ge, M., Zhang, S., Li, Z., Zhang, X., Francisco, J. S., Zeng,
691 X. C., 2018. Self-catalytic reaction of SO₃ and NH₃ to produce sulfamic acid and its implication to atmospheric
692 particle formation. *J. Am. Chem. Soc.*, 140, 11020-11028, <http://doi.org/10.1021/jacs.8b04928>

693 Li, J., Han, Z., Yao, X., Xie, Z., Tan, S., 2019. The distributions and direct radiative effects of marine aerosols over
694 East Asia in springtime. *Sci. Total Environ.*, 651, 1913-1925, <http://doi.org/10.1016/j.scitotenv.2018.09.368>

695 Li, L., Hoffmann, M. R., Colussi, A. J., 2018. Role of nitrogen dioxide in the production of sulfate during Chinese

696 haze-aerosol episodes. *Environ. Sci. Technol.*, 52, 2686-2693, <http://doi.org/10.1021/acs.est.7b05222>

697 Li, W., McKee, M. L., 1997. Theoretical study of OH and H₂O addition to SO₂. *J. Phys. Chem. A*, 101, 9778-9782,
698 <http://doi.org/10.1021/jp972389r>

699 Liu, J., Fang, S., Wang, Z., Yi, W., Tao, F., Liu, J., 2015. Hydrolysis of sulfur dioxide in small clusters of sulfuric
700 acid: mechanistic and kinetic study. *Environ. Sci. Technol.*, 49, 13112-13120,
701 <http://doi.org/10.1021/acs.est.5b02977>

702 Liu, T., Abbatt, J. P. D., 2020. An experimental assessment of the importance of S(IV) oxidation by hypohalous
703 acids in the marine atmosphere. *Geophys. Res. Lett.*, 47, 1-8, <http://doi.org/10.1029/2019GL086465>

704 Lu, T., Chen, F., 2012. Multiwfn: a multifunctional wavefunction analyzer. *J. Comput. Chem.*, 33, 580-592,
705 <http://doi.org/10.1002/jcc.22885>

706 Manzetti, S., Lu, T., 2013. The geometry and electronic structure of aristolochic acid: possible implications for a
707 frozen resonance. *J. Phys. Org. Chem.*, 26, 473-483, <http://doi.org/10.1002/poc.3111>

708 Mardirossian, N., Head-Gordon, M., 2016. How accurate are the minnesota density functionals for noncovalent
709 interactions, isomerization energies, thermochemistry, and barrier heights involving molecules composed of
710 main-group elements? *J. Chem. Theory Comput.*, 12, 4303-4325, <http://doi.org/10.1021/acs.jctc.6b00637>

711 Marenich, A. V., Cramer, C. J., Truhlar, D. G., 2009. Performance of SM6, SM8, and SMD on the SAMPL 1 test
712 set for the prediction of small-molecule solvation free energies. *J. Phys. Chem. B*, 113, 4538-4543,
713 <http://doi.org/10.1021/jp809094y>

714 McArdle, J. V., Hoffmann, M. R., 1983. Kinetics and mechanism of the oxidation of aquated sulfur dioxide by
715 hydrogen peroxide at low pH. *J. Phys. Chem.*, 87, 5425-5429, <http://doi.org/10.1021/j150644a024>

716 Meijer, E. J., Sprik, M., 1998. A density functional study of the addition of water to SO₃ in the gas phase and in
717 aqueous solution. *J. Phys. Chem. A*, 102, 2893-2898, <http://doi.org/10.1021/jp972146z>

718 Miguel, E. L. M., Santos, C. I. L., Silva, C. M., Jr. Pliego, J. R., 2016. How accurate is the SMD model for predicting
719 free energy barriers for nucleophilic substitution reactions in polar protic and dipolar aprotic solvents? *J. Braz.*
720 *Chem. Soc.*, 27, 2055-2061, <http://doi.org/10.5935/0103-5053.20160095>

721 Minakata, D., Song, W., Mezyk, S. P., Cooper, W. J., 2015. Experimental and theoretical studies on aqueous-phase
722 reactivity of hydroxyl radicals with multiple carboxylated and hydroxylated benzene compounds. *Phys. Chem.*
723 *Chem. Phys.*, 17, 11796-11812, <http://doi.org/10.1039/c5cp00861a>

724 Murray, J. S., Politzer, P., 2011. The electrostatic potential: an overview. *Wires. Comput. Mol. Sci.*, 1, 153-163,
725 <http://doi.org/10.1002/wcms.19>

726 Neese, F., 2012. The ORCA program system. *Wires. Comput. Mol. Sci.*, 2, 73-78, <http://doi.org/10.1002/wcms.81>

727 Novakov, T., Rivera-Carpio, C., Penner, J. E., Rogers, C. F., 1994. The effect of anthropogenic sulfate aerosols on
728 marine cloud droplet concentrations. *Tellus*, 46, 132-141, <http://doi.org/10.1034/j.1600-0889.1994.t01-1->
729 00005.x

730 O'Dowd, C. D., de Leeuw, G., 2007. Marine aerosol production: a review of the current knowledge. *Phil. Trans. R.*
731 *Soc. A.*, 365, 1753-1774, <http://doi.org/10.1098/rsta.2007.2043>

732 Ostovari, H., Zahedi, E., Sarvi, I., Shiroudi, A., 2018. Kinetic and mechanistic insight into the formation of
733 amphetamine using the Leuckart–Wallach reaction and interaction of the drug with GpCCpG base-pair step of
734 DNA: a DFT study. *Monatshefte für Chemie - Chemical Monthly*, 149, 1045-1057,
735 <http://doi.org/10.1007/s00706-018-2145-7>

736 Peterson, K. A., Adler, T. B., Werner, H., 2008. Systematically convergent basis sets for explicitly correlated
737 wavefunctions: The atoms H, He, B–Ne, and Al–Ar. *J. Chem. Phys.*, 128, 84102,
738 <http://doi.org/10.1063/1.2831537>

739 Pszenny, A. A. P., Moldanová, J., Keene, W. C., Sander, R., Maben, J. R., Martinez, M., Crutzen, P. J., Perner, D.,
740 Prinn, R. G., 2003. Halogen cycling and aerosol pH in the Hawaiian marine boundary layer. *Atmos. Chem.*
741 *Phys.*, 3, 4701-4753, <http://doi.org/10.5194/acpd-3-4701-2003>

742 Rong, H., Liu, L., Liu, J., Zhang, X., 2020. Glyoxylic sulfuric anhydride from the gas-phase reaction between
743 glyoxylic acid and SO₃: A potential nucleation precursor. *J. Phys. Chem. A*, 124, 3261-3268,
744 <http://doi.org/10.1021/acs.jpca.0c01558>

745 Roussel, M. R., 1985. Diffusion-influenced reactions. University of Lethbridge.

746 Shaka, H., Robertson, W. H., Finlayson-Pitts, B. J., 2007. A new approach to studying aqueous reactions using
747 diffuse reflectance infrared Fourier transform spectrometry: application to the uptake and oxidation of SO₂ on
748 OH-processed model sea salt aerosol. *Phys. Chem. Chem. Phys.*, 9, 1980-1990,
749 <http://doi.org/10.1039/b612624c>

750 Shen, X., Lee, T., Guo, J., Wang, X., Li, P., Xu, P., Wang, Y., Ren, Y., Wang, W., Wang, T., Li, Y., Carn, S. A.,
751 Collett, J. L., 2012. Aqueous phase sulfate production in clouds in eastern China. *Atmos. Environ.*, 62, 502-
752 511, <http://doi.org/10.1016/j.atmosenv.2012.07.079>

753 Smoluchowski, M. V., 1917. Mathematical Theory of the kinetics of the coagulation of colloidal solutions. *Z. Phys.*
754 *Chem.*, 92, 129-168,

755 Stokes, G. G., 1903. Mathematical and physical papers. U.K. Cambridge University Press.

756 Torrent-Sucarrat, M., Francisco, J. S., Anglada, J. M., 2012. Sulfuric acid as autocatalyst in the formation of sulfuric
757 acid. *J. Am. Chem. Soc.*, 134, 20632-20644, <http://doi.org/10.1021/ja307523b>

758 Truhlar, D. G., Hase, W. L., Hynes, J. T., 1983. Current status of transition-state theory. *J. Phys. Chem.*, 87, 2664-
759 2682, <http://doi.org/10.1021/j100238a003>

760 Truhlar, D. G., 1985. Nearly encounter-controlled reactions: The equivalence of the steady-state and diffusional
761 viewpoints. *J. Chem. Educ.*, 104-106, <http://doi.org/10.1021/ed062p104>

762 Voegele, A. F., Loerting, T., Tautermann, C. S., Hallbrucker, A., Mayer, E., Liedl, K. R., 2004. Sulfurous acid
763 (H_2SO_3) on Io. *Icarus*, 169, 242-249, <http://doi.org/10.1016/j.icarus.2003.11.012>

764 Voegele, A. F., Tautermann, C. S., Loerting, T., Hallbrucker, A., Mayer, E., Liedl, K. R., 2002. About the stability
765 of sulfurous acid (H_2SO_3) and its dimer. *Chem. Eur. J.*, 8, 5644-5651, <http://doi.org/10.1002/chin.200313001>

766 Vogt, R., Crutzen, P. J., Sander, R., 1996. A mechanism for halogen release from sea-salt aerosol in the remote
767 marine boundary layer. *Nature*, 383, 327-330, http://doi.org/10.1007/978-3-319-27460-7_8

768 von Glasow, R., Sander, R., Bott, A., Crutzen, P. J., 2002. Modeling halogen chemistry in the marine boundary layer
769 2. Interactions with sulfur and the cloud-covered MBL. *J. Geophys. Res. Atmos.*, 107, 1-2,
770 <http://doi.org/10.1029/2001JD000943>

771 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y., Peng, J., Guo, S., Meng, J., Li,
772 J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., Cao, J., An, Z., Zhou, W., Li, G., Wang, J., Tian, P., Marrero-
773 Ortiz, W., Secrest, J., Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., Wang, W., Huang, Y., Wang, Y., Zhu,
774 Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb,
775 C. E., Molina, M. J., 2016. Persistent sulfate formation from London Fog to Chinese haze. *P. Natl. Acad. Sci.*
776 *Usa.*, 113, 13630-13635, <http://doi.org/10.1073/pnas.1616540113>

777 Wang, J., Li, J., Ye, J., Zhao, J., Wu, Y., Hu, J., Liu, D., Nie, D., Shen, F., Huang, X., Huang, D. D., Ji, D., Sun, X.,
778 Xu, W., Guo, J., Song, S., Qin, Y., Liu, P., Turner, J. R., Lee, H. C., Hwang, S., Liao, H., Martin, S. T., Zhang,
779 Q., Chen, M., Sun, Y., Ge, X., Jacob, D. J., 2020. Fast sulfate formation from oxidation of SO_2 by NO_2 and
780 HONO observed in Beijing haze. *Nat. Commun.*, 11, 1-7, <http://doi.org/10.1038/s41467-020-16683-x>

781 Wang, Y., Zhang, Q., Jiang, J., Zhou, W., Wang, B., He, K., Duan, F., Zhang, Q., Philip, S., Xie, Y., 2014. Enhanced
782 sulfate formation during China's severe winter haze episode in January 2013 missing from current models. *J.*
783 *Geophys. Res. Atmos.*, 119, 10425-10440, <http://doi.org/10.1002/2013JD021426>

784 Wang, Y. H., Liu, Z. R., Zhang, J. K., Hu, B., Ji, D. S., Yu, Y. C., Wang, Y. S., 2015. Aerosol physicochemical
785 properties and implications for visibility during an intense haze episode during winter in Beijing. *Atmos. Chem.*

786 Phys., 15, 3205-3215, <http://doi.org/10.5194/acp-15-3205-2015>

787 Xu, L., Coote, M. L., 2019. Methods to improve the calculations of SMD solvation free energies and associated
788 aqueous pKa values: comparison between choosing an optimal theoretical level, solute cavity scaling and using
789 explicit solvent molecules. *J. Phys. Chem. A*, 123, 7430-7438, <http://doi.org/10.1021/acs.jpca.9b04920>

790 Yan, J., Jung, J., Lin, Q., Zhang, M., Xu, S., Zhao, S., 2020. Effect of sea ice retreat on marine aerosol emissions in
791 the Southern Ocean, Antarctica. *Sci. Total Environ.*, 745, 140773,
792 <http://doi.org/10.1016/j.scitotenv.2020.140773>

793 Zhang, F., Wang, Y., Peng, J., Chen, L., Sun, Y., Duan, L., Ge, X., Li, Y., Zhao, J., Liu, C., Zhang, X., Zhang, G.,
794 Pan, Y., Wang, Y., Zhang, A. L., Ji, Y., Wang, G., Hu, M., Molina, M. J., Zhang, R., 2020. An unexpected
795 catalyst dominates formation and radiative forcing of regional haze. *P. Natl. Acad. Sci. Usa.*, 117, 3960-3966,
796 <http://doi.org/10.1073/pnas.1919343117>

797 Zhang, H., Wang, W., Pi, S., Liu, L., Li, H., Chen, Y., Zhang, Y., Zhang, X., Li, Z., 2018. Gas phase transformation
798 from organic acid to organic sulfuric anhydride: possibility and atmospheric fate in the initial new particle
799 formation. *Chemosphere*, 212, 504-512, <http://doi.org/10.1016/j.chemosphere.2018.08.074>

800 Zhang, Y., Bao, F., Li, M., Chen, C., Zhao, J., 2019. Nitrate-enhanced oxidation of SO₂ on mineral dust: a vital role
801 of a proton. *Environ. Sci. Technol.*, 53, 10139-10145, <http://doi.org/10.1021/acs.est.9b01921>

802 Zhao, J., Kong, X., He, K., Xu, H., Mu, J., 2020. Assessment of the radiation effect of aerosols on maize production
803 in China. *Sci. Total Environ.*, 720, 137567, <http://doi.org/10.1016/j.scitotenv.2020.137567>

804 Zhao, Y., Truhlar, D. G., 2008. The M06 suite of density functionals for main group thermochemistry,
805 thermochemical kinetics, noncovalent interactions, excited states, and transition elements: two new functionals
806 and systematic testing of four M06-class functionals and 12 other functionals. *Theor. Chem. Acc.*, 120, 215-
807 241, <http://doi.org/10.1007/s00214-007-0310-x>

808 Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., Kimoto, T., 2015.
809 Heterogeneous chemistry: a mechanism missing in current models to explain secondary inorganic aerosol
810 formation during the January 2013 haze episode in North China. *Atmos. Chem. Phys.*, 15, 2031-2049,
811 <http://doi.org/10.5194/acp-15-2031-2015>

812