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Sources and dynamics of submicron aerosol during the autumn onset of the air pollution season in Delhi, India

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Abstract. New Delhi, India is the most polluted megacity in the world and routinely experiences high particulate matter (PM) concentrations. As part of the Delhi Aerosol Supersite Study, we have been measuring PM₁ concentration and composition in Delhi near continuously since January 2017. This manuscript focuses on autumn, one of the most

polluted seasons in Delhi when PM_1 concentrations steadily increase throughout the season and can exceed $1000 \mu g m^{-3}$ during episodic events. Positive matrix factorization on the organic aerosol (OA) spectrum suggests comparable seasonal average contributions from HOA (Hydrocarbon-like OA), BBOA (Biomass-Burning OA) and OOA (Oxidized- OA), with BBOA dominating during episodic events. We demonstrate the influence of regional sources such as agricultural burning during this season through temporal trends of pollutants, PMF factors, meteorology, and non-parametric wind regression analysis. We use inorganic fragment ratios to show the influence of metals during the festival of Diwali. Furthermore, we demonstrate the influence of transitioning meteorology in governing PM_1 composition through the season. Overall, our analysis provides novel insights into the factors controlling PM_1 during one of the most polluted seasons in Delhi.

1. Introduction

Exposure to fine particulate matter (PM) is a major risk factor for cardiovascular, respiratory, and other diseases¹⁻⁵. Globally, outdoor $PM_{2.5}$ (PM of size $< 2.5 \mu m$) resulted in 4.2 million premature deaths in 2016⁶, predominantly in low- and middle-income countries⁷. In India, exposure to $PM_{2.5}$ has been estimated to reduce the average life expectancy by ~ 1.5 years, compared to ~ 0.4 years in the U.S.⁸. Delhi, India is the second most populated megacity in the world (current population 28 million) and routinely experiences high air pollution levels, thereby making it a high-risk location for exposure to PM and associated adverse health effects^{7,9-11}.

Atmospheric PM can originate from a range of natural and anthropogenic sources¹²⁻¹³. Primary aerosols are directly emitted, whereas secondary aerosols are formed from gaseous precursors via atmospheric oxidation and subsequent partitioning to the aerosol phase¹³⁻¹⁵. The atmospheric fate and concentration of ambient PM depends on numerous factors including primary emissions, meteorological factors (wind speed, direction, temperature, relative humidity, planetary boundary layer height), and the atmospheric oxidizing capacity¹⁶⁻²³.

Local sources of primary PM in Delhi include transportation, domestic biomass and trash burning, cooking, industrial and construction activities²⁴⁻³⁰. Delhi is also downwind of many agricultural states such as Punjab and Haryana which can be a source of PM from agricultural burning³¹⁻³⁵. Further, Delhi experiences cool winters with shallow boundary layer heights and frequent temperature inversions, which trap pollutants within the boundary layer causing especially polluted conditions³⁶⁻³⁹. Rapid photochemical processing in Delhi contributes to the importance of secondary pollutants³¹.

Autumn (mid-Sep - Nov) is an especially important transitional period in air quality in Delhi and across North India. During autumn, PM concentrations rapidly increase from their annual monsoon minima (Jun-Aug, typical levels 30-50 $\mu\text{g m}^{-3}$ $\text{PM}_{2.5}$) to reach daily levels above 200 $\mu\text{g m}^{-3}$ $\text{PM}_{2.5}$ and episodic high concentrations in excess of 500-1000 $\mu\text{g m}^{-3}$. While the sources, atmospheric dynamics, and chemistry of aerosols in Delhi have received sustained attention in recent years, we provide here new perspectives with a detailed

investigation of the mechanisms driving the rapid increase in particle-phase pollution during autumn months.

Several features of the ambient aerosol in Delhi during autumn months merit brief introduction. (check these refs: ^{32,35}). A few important seasonal sources of pollution are especially active during autumn months: a few weeks of intensive burning of rice (“paddy”) crop stubble in the upwind states of Punjab and Haryana^{33,40}; the festivals of Dussehra and especially Diwali, which involve the burning of fireworks⁴¹⁻⁴²; and the resumption of brick-making season in uncontrolled brick kilns after the end of monsoon rains. Further, this season observes a sharp transition in meteorology since it is between monsoon (Jul – mid Sep), which is hot, humid, and has higher planetary boundary layer height and mostly southern winds, and winter (Dec - mid Feb), which is cold and has lower planetary boundary layer height and mostly northwestern winds³². Transitioning meteorology likely contributes to variability in sources impacting the city (due to changing wind direction) and also changes particle concentrations due to thermodynamic (partitioning) effects. As shown in Fig.S1, a rapid increase of $\sim 20\text{-}30\text{ }\mu\text{g m}^{-3}$ / week in autumn months is typical over the recent years of the the observational record. In addition, large episodic spikes in pollution – sometimes but not always coincident with festival periods – are common. Here, we focus on autumn 2018, a representative recent example of autumn air pollution for which we made intensive measurements of chemical composition. As part of the Delhi Aerosol Supersite study we have been measuring the submicron aerosol composition, concentration and size distribution using a suite of online instrumentation since January 2017.

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The main objectives of our analyses are to understand factors contributing to the high concentrations observed in autumn 2018 by 1) interpreting the temporal trends in sub-micron composition and concentration, 2) using positive matrix factorization (PMF) to understand the nature and sources of organic aerosol (OA) and 3) interpreting the influence of meteorological parameters.

2. Methods

2.1 Sampling site, instrumentation, and data processing

The Delhi Aerosol Supersite is located at the Indian Institute of Technology Delhi (IITD) campus in South Delhi. Details on the sampling site, instrumentation and set up can be found in Gani et al., 2019³². Briefly, the bulk composition of non-refractory PM₁ (NR-PM₁) is measured using an Aerodyne Aerosol Chemical Speciation Monitor (ACSM, Aerodyne Research, Billerica MA)⁴³. Details on ACSM calibration are presented in section S1 of the supplementary information (SI). Black carbon (BC, at 880 nm), ultraviolet absorbing particulate matter (UV-PM, at 370 nm) and their difference (deltaC) are measured using a multi-channel aethalometer (Magee Scientific Model AE33, Berkeley, CA)⁴⁴. Particle size distributions (PSD) are measured using a scanning mobility particle sizer (SMPS, TSI, Shoreview, MN). The SMPS was not operational during the Diwali period from Nov 04th-8th.

The SMPS scanned from 12 to 560 nm, and we used a mode fitting algorithm⁴⁵ in the mass domain to estimate the PSD between 560 and 1000 nm³². Although the ACSM recorded data every ~1 minute, it was post averaged to 15 points (~15 minutes) for performing PMF and 60 minutes for comparison with SMPS data. Further details on ACSM data processing are

presented in section S1. The biomass burning and fossil fuel fractions of BC (BC_{bb} and BC_{ff} respectively) were computed using the model of Sandradewi et al⁴⁶.

2.2 Source apportionment

Positive matrix factorization (PMF) is a receptor modeling tool which has been widely used in combination with ambient aerosol mass spectrometry data to apportion “factors” that serve as proxies for various PM sources and types⁴⁷⁻⁵⁶. The factors resolved usually correspond to emissions from primary sources including hydrocarbon like OA, “HOA”, used as a proxy for traffic emissions and biomass burning OA, “BBOA”, used as a proxy for biomass burning emissions. OA formed from secondary reactions, secondary organic aerosol (SOA) is typically resolved as oxidized OA, “OOA”^{19,31,47}. Further separation of factors corresponding to sources such as cooking, coal combustion and trash burning may be possible depending on their contribution to total OA, correlation to other sources, as well as the mass resolution of the instrument^{25,47,48,57-60}.

We performed PMF on OA mass spectral data from the ACSM, utilizing the PMF Evaluation Tool, PET⁴⁷ which uses the PMF2 algorithm⁶¹. We selected m/z 12 to 120 in the OA mass spectra due to a low signal to noise ratio at higher m/z ^{50,62}. We assigned physical meaning to the factors based on the abundance of specific m/z in their spectra, correlation with reference spectra^{47,63} as well as with the time series of external tracers (e.g. CO, NO_x, BC_{ff} , BC_{bb} , deltaC) where possible⁴⁷. A three-factor solution was chosen to best represent the data as discussed further in section 3.2. A value of zero was chosen for SEED and FPEAK because non-zero values either had no significant effect on the solution or led to unreasonable factors

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due to factor splitting. Further details on the PMF runs and number of factors, SEED and FPEAK scenarios are presented in section S2 (and Table S1, Fig.S2).

2.3 Other data

We retrieved data on hourly wind speed, wind direction (10 m above ground level), and planetary boundary layer height (H) from the NASA meteorological reanalysis dataset MERRA2⁶⁴, and data on temperature and relative humidity from the Indira Gandhi International Airport (~ 8 kilometers west of our site). We obtained daily fire counts in north India from the NASA fire information for resource management system (FIRMS) using the moderate resolution imaging spectroradiometer (MODIS, collection 6) data⁶⁵⁻⁶⁷, CO, NO_x and PM_{2.5} data from the Central Pollution Control Board (CPCB) Central Control Room web interface (CCR) with additional processing steps summarized in section S1.

We performed non-parametric wind regression analysis using ZeFir v3.7⁶⁸ in order to identify potential source origins of the PM₁ species and the PMF factors. Briefly speaking, this is a source-to-receptor apportionment model which uses non-parametric kernel smoothing methods to apportion the observed concentrations to sectors defined by wind speed and wind direction⁶⁹. In order words, it estimates the concentration measured by the instrument as a function of wind speed and direction (the sources sampled depend on wind direction). Further details about the method are presented in section S3.

3. Results and Discussion

3.1 Temporal trends in PM_1 and meteorological parameters

The average C- PM_1 (NR- PM_1 + BC) concentration gradually increased from $\sim 65 \mu g m^{-3}$ (averaged for the last two weeks) in Sep to $\sim 220 \mu g m^{-3}$ in Nov (averaged for the whole month). The maximum hourly concentration of the study ($\sim 1000 \mu g m^{-3}$) was observed during the early morning hours ($\sim 06:00$ a.m.) of Nov 5th (Fig.2). The C- PM_1 concentration averaged for the entire season was $161 \mu g m^{-3}$ (Table S2), higher than the concentrations observed in spring, summer and monsoon and $\sim 20\%$ lower than that observed in winter (diurnal variation of average concentration/composition across all seasons shown in Fig.S3)³², which experiences lower temperatures, boundary layer height and likely more local emissions from domestic heating³⁰.

As shown in Fig.1a (weekly moving average), the overall increase in C- PM_1 through the season was accompanied by a decrease in average ventilation coefficient (VC = planetary boundary layer height $\times$ wind speed) and temperature (T). Further, concentrations of NO_x and CO (which are mostly associated with local sources) varied with $1/VC$ while C- PM_1 increased consistently and rapidly until mid Nov, indicating the impact of additional sources such as regional PM.

Figure 1: Weekly moving average of a) C-PM₁ (= NR-PM₁ + BC) , CO, NO_x , temperature and inverse of ventilation coefficient (VC = planetary boundary layer height × wind speed) b) wind speed (WS, colors represent the wind direction) c) NR-PM₁ species concentrations d) NR-PM₁ composition. e) average diurnal variation of C-PM₁ species and f) average C-PM₁ composition in autumn 2018.

Fig.1c and d show the weekly moving averages of the NR-PM₁ species concentrations and composition, respectively through late monsoon, autumn and early winter. Sulfate (SO₄²⁻) concentration was initially higher than the other inorganic species, but was outcompeted by nitrate (NO₃⁻), ammonium (NH₄⁺) by Nov. Since SO₄²⁻ is non-volatile, its concentration is not expected to be influenced by changes in temperature. On the other hand, NO₃⁻ is volatile and partitions to the particle phase at lower temperatures, as ammonium nitrate (NH₄NO₃). The increase in chloride (Cl⁻) concentration was more significant in (early) winter than autumn,

consistent with further reduction in temperature and the volatile nature of ammonium chloride (NH_4Cl)⁷⁰. Thus, the increase in NH_4^+ , NO_3^- and Cl^- concentrations in autumn was likely driven by VC and temperature while the increase in SO_4^{2-} concentration was driven mainly by VC. Further, the largest increase in concentration in autumn was observed in organics, pointing to the role of increased diversity of sources during this season. Overall, SO_4^{2-} fraction decreased, organics and NO_3^- fractions increased through autumn (Fig.1d). In early winter (Dec 1st-31st), the organics fraction decreased as NO_3^- and Cl^- fractions increased further, reflective of differences in composition in autumn versus winter.

The average organics fraction in C-PM_1 in autumn 2018 was ~64% (Fig.1f), higher than other seasons (where it was less than or equal to ~55%)³², indicating that the source mixture in autumn is different than in other seasons (Fig.S3). As shown in Fig.1e, OA peaked during the morning hours between 07:00 – 09:00 a.m. and during the nighttime hours between 09:00 - 11:00 p.m., due to the influence of local emissions (such as traffic and local biomass burning). Interestingly, unlike other seasons which observed a dip in OA concentration during the early morning hours (Fig.S3), it remained relatively constant through 00:00 - 7:00 a.m., suggestive of the influence of additional sources during those hours.

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Figure 2: High concentration episode from Nov 04th -08th 2018 (Hourly averages are shown)

Overall, the hourly averaged C-PM₁ data compared well with the SMPS PM₁ estimate (Fig.S4; $R^2 = 0.91$). It also compared reasonably well with the regulatory monitor-based PM_{2.5} data (Fig.S5) on most days except the period from Nov 7th-9th. On the night of Diwali (Nov 7th), the maximum C-PM₁ concentration ($\sim 230 \mu\text{gm}^{-3}$) was not as high as the concentrations observed on the earlier days (Fig.2). Although the PM_{2.5} monitoring sites recorded peaks on/after the Diwali day, there were differences in the peak concentration values as well as the time at which they were observed (see temporal trend in Fig.S6b). Meteorological data during the period indicate a change in wind direction from North West to North East/South East and slower winds starting Nov 7th (Fig.S6a). Thus, the lower concentration measured on the night of Diwali might have been due to the instrument sampling a different local plume as a result of the relatively stagnant wind conditions and the absence of high emission sources in the proximity of our site located inside the IIT-Delhi campus. Further, fireworks emit mineral dust and metals which are not detected by the ACSM⁷¹. We observed some evidence of metal compounds during Diwali, as discussed in section 3.3. Furthermore, previous studies have found that a large fraction of PM from fireworks may be of size greater

than $1\mu\text{m}$ ⁷¹, which would not have been sampled by the ACSM. Additionally, there was an influence of agricultural burning on the earlier days considering that the winds were faster and from the north-west direction – a topic we discuss further in sections 3.2 and 3.4. Thus, we did not observe such high concentrations on Diwali because of meteorology, as well as the limitations of the ACSM in measuring metals, dust, and particles of size greater than $1\mu\text{m}$.

The average particle number (PN) concentration in autumn 2018 (39000 cm^{-3} ; Table S2) was comparable to that observed in autumn 2017 (38000 cm^{-3})⁷². Further, the diurnal variation in particle number and mass concentrations (Fig.S7) of nucleation (sub-25 nm), Aitken (25-100 nm), and accumulation (100-1000 nm) modes compared remarkably well between the two periods, indicating little inter-annual variation in the sources and processes governing these modes, and suggesting that our data for autumn 2018 is representative of autumns in Delhi.

3.2 Insights from source apportionment

PMF performed on OA for this season resolved three factors – two primary factors, HOA, BBOA and a secondary factor, OOA, which were identified based on their correlation with reference factors (Pearson $R > 0.9$), and external tracers (Figs.S8-9). Intense and sporadic events (e.g. fireworks, intense wood burning etc.) can influence the PMF results and are therefore usually excluded when running the model⁷³⁻⁷⁴. We tested the influence of high concentration episodes observed from Nov 4th-8th on our results by running a PMF model with the exclusion of that period. The resulting factors' mass spectra and time series were highly correlated with the PMF results obtained for the whole period (Pearson $R = 0.99$;

Figs.S8-9) and the slopes of the scatter plots of the respective PMF factor concentrations were close to 1 (Fig.S10). Thus, these high concentration episodes had little impact on the overall results, and we therefore use results from PMF performed on the whole period for the discussion below.

3.2.1 Factor mass spectra

The mass spectra (MS) of HOA has increased signals at m/z 41,43,55,57,69,71 etc., usually corresponding to the fragments of aliphatic hydrocarbons (C_nH_{2n-1} and C_nH_{2n+1}) resulting from traffic emissions (Fig.S11a). It is highly correlated to the HOA MS obtained in spring 2018 (Pearson $R = 0.99$; Fig.S12a), suggesting a consistent signature of this source (BBOA and HOA separation was achieved only in spring 2018 of our previous analysis³¹). A relatively high ratio of m/z 55/57 (~ 1.23 , compared to ~ 1.05 in reference HOA,) suggests potential influence of cooking organic aerosol (COA) on this factor^{53,58,62,75-77}. However, the enhancement is lower than that in the reference COA factor (~ 2.92 ; Fig.S12b), indicating that it was a minor contributor. This is also consistent with the plot of f_{55} (fraction of OA mass at m/z 55) versus f_{57} in a triangle plot⁷⁸ (Fig.S12c), where most of the points in autumn 2018 lie close to the HOA line(s), indicative of minor contribution from COA. The time series (TS) of HOA is correlated to the external tracers for traffic emissions, CO, NO_x , BC and BC_{ff} (Fig.S8a).

BBOA MS has enhanced signals at m/z 29, 60 and 73 (Fig.S11b), usually associated with the fragments CHO^+ , $C_2H_4O_2^+$ and $C_3H_5O_2^+$, respectively, resulting from anhydrous sugars such as levoglucosan^{57,79,80}. Its MS is correlated to the BBOA MS obtained in spring 2018 (Pearson

$R = 0.9$) with some differences including a higher f_{44} and f_{60} in autumn (Fig.S13). Fig.S14a depicts f_{44} versus f_{43} for autumn BBOA, its comparison with reference BBOA, as well as the profiles of fresh and aged (~ 5 hours) rice/wheat OA from literature⁸¹. The data points for autumn 2018 BBOA lie in between fresh and aged agricultural burning OA, indicating that it likely had contribution from regional (aged) as well as local (fresh) biomass burning sources. Further, Fig.S14b shows that autumn 2018 BBOA has a higher f_{60} and f_{44} than the reference BBOA. Differences in f_{60} may be due to increased contribution from sources with higher levoglucosan content⁸², while a higher f_{44} points towards faster aging and/or contribution from regional sources⁸¹. The autumn BBOA TS is correlated to the external tracers for biomass burning such as $\delta^{13}C$ and BC_{bb} (Fig.S8b). It is also correlated to chloride (Pearson $R = 0.95$), indicating that the chloride during this season was likely associated with biomass burning (also see sections 3.3 and 3.4). We compared the average OA spectra of autumn and winter to get an understanding of the overall OA source mixture during the two periods since we were not able to resolve BBOA in the winter seasons previously analysed³¹. Autumn OA is more oxidized than winter OA (higher f_{44} , Fig.S15), pointing towards the role of increased photo-chemical processing and/or contribution from regionally transported PM. Regional contribution is also suggested by the aethalometer data - high $\delta^{13}C$, and $BB_{percent}$ values were observed in autumn (peaking in Nov) and winter, indicating contribution from biomass burning during both the periods (see temporal trend in Fig.S16a). However, the compensation parameter values (Fig.S16b; used to correct for spot loading effects), which are expected to be higher for fresher aerosols (close to the source) and lower for processed aerosols (away from the source)⁴⁴ were lower during the autumn period, indicating increased influence of regional biomass burning sources on the $\delta^{13}C$ observed during this

season. This suggests that the BBOA in autumn has a higher contribution from regional agricultural biomass burning whereas the BBOA in winter has a relatively higher contribution from domestic biomass burning.

The f_{44} signal in OOA (Fig.S11c) is higher than the corresponding signal in the reference OOA spectrum (by a factor of ~ 1.4 ; Fig.S17), similar to other seasons, suggestive of rapid photochemical processing of fresh emissions³¹.

3.2.2 Temporal trends of PMF factors

Figure 3: a) Weekly moving average of PMF factor concentrations. The inverse of ventilation coefficient (VC=planetary boundary layer height $\times$ wind speed) has also been added to track the change in meteorology. b) Weekly moving average of PMF factor fractions. The pie chart depicts the average OA composition. Average diurnal variations in factor concentrations and fractions are depicted in figures c) and d) respectively.

Concentrations of all factors increased through the season (Fig.3a), but the fractional increase in BBOA concentration was largest, increasing not only with time (Fig.3b) but also with the concentration of NR-PM₁ (Fig.S18b). The clear peak in BBOA concentrations during

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3 the first week of Nov suggests that the episodic events Nov 5th and 6th were due to increased
4 biomass burning (15-minute averaged data shown in Fig.S18a). Furthermore, the variation
5 in HOA concentration, (which is associated with local sources such as traffic) anti-correlated
6 with the variation in ventilation coefficient, indicating that its increase was mostly associated
7 with changing meteorology. In contrast, the increase in BBOA, especially between mid-Oct and
8 mid-Nov did not correlate strongly with the decrease in ventilation coefficient, pointing
9 towards the influence of other sources. This increase in BBOA concentration was consistent
10 with the increase in fire counts in northern India during the period (temporal trend of daily
11 fire counts in Fig.S19), indicating that regional agricultural fires contributed to BBOA during
12 this period. The BBOA fraction remained high even when the fire counts dropped towards
13 the end of Nov, presumably due to an increase in the contribution of BBOA sources such as
14 domestic heating owing to the lower ambient temperatures (Fig.2).

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17 In terms of diurnal variation in concentration (Fig.3c), HOA had morning (07:00-09:00 a.m.)
18 and night time (09:00-11:00 p.m.) peaks corresponding to traffic hours, with the night time
19 peak being larger than the day time peak, likely due to the absence of photochemistry (which
20 converts HOA to OOA) and some influence of cooking activities during those hours (section
21 3.2.1). The dip in the primary OA (HOA and BBOA) factor concentrations during the
22 afternoon hours was concurrent with the increase in planetary boundary layer height,
23 temperature (see their diurnal variation in Fig.S20) and photochemistry. BBOA
24 concentration increased through the nighttime and early morning hours after the afternoon
25 dip, indicative of BBOA source(s) during those hours, such as agricultural burning. The
26 secondary factor, oxidized OA (OOA), observed a small day time peak (08:00 a.m. – 10:00
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a.m.), but otherwise remained relatively constant though the day, consistent with photochemical formation mostly off-setting the effects of the increasing PBLH during the afternoon.

In terms of diurnal variation in OA composition (Fig.3d), BBOA dominated during the early morning (02:00-08:00 a.m), constituting ~ 37-43% of OA , while HOA dominated during the night (09:00 p.m. - 01:00 a.m.), constituting ~39-41% of OA. OOA dominated during the day (09:00 a.m. - 08:00 p.m), constituting 41-65% of OA, consistent with daytime photochemistry being the main pathway for formation of oxygenated OA. In the triangle plot⁴² of diurnally averaged f_{44} versus f_{43} (Fig.S21b), all the data points lie within the triangle, with the afternoon data points occupying the top left position. On average, all the three factors had comparable contribution to OA (HOA – 30%, BBOA-33%, and OOA-37%), suggesting similar importance of biomass burning (BBOA), traffic (HOA) and secondary OA (OOA) on average during this season (Fig.3b). The statistical metrics – standard deviation, arithmetic mean-median and geometric standard deviation, which indicate the influence of outliers (i.e., episodic events) are highest for BBOA (Table S3), due to the influence of agricultural burning.

3.3 Insights from inorganic fragment ratios

The enhancement in the ratio of the NO_3^- fragments at m/z 30 (NO^+) to m/z 46 (NO_2^+) with respect to the value obtained for pure ammonium nitrate points towards the relative importance of organonitrates⁸³. During the autumn season, the ratio decreased through the progression of the season as total particulate NO_3^- increased (see temporal trend in

Fig.S22a), indicating that the relative contribution of organonitrates decreased through the season, likely a result of an increase in the inorganic nitrate fraction at lower temperatures (Fig.2). Interestingly, the $\text{NO}^+/\text{NO}_2^+$ ratio started to increase on the night of Nov 7th (Diwali) and reached $\sim 3 \times$ calibration value (of that of ammonium nitrate) on the early hours of the following day (Fig.S22b), likely due to the presence of metal nitrates such as sodium nitrate (NaNO_3) and potassium nitrate (KNO_3) in fireworks⁸⁴⁻⁸⁵ which have been shown to produce high $\text{NO}^+/\text{NO}_2^+$ ratio⁸⁶. Further, similar to the enhancement in $\text{NO}^+/\text{NO}_2^+$, an enhancement in $\text{SO}^+/\text{H}_y\text{SO}_x^+$ and $\text{SO}_2^+/\text{H}_y\text{SO}_x^+$ was also observed (see temporal trend in Fig.S23) on the night of Diwali, likely due to the presence of metal sulfates in the fireworks^{85,86,87}.

Quantification of potassium (K^+ , m/z 39) using ACSM is challenging because of uncertainties in vaporization and surface ionization as well as interference of the organic fragment C_3H_3^+ at m/z 39⁸⁸. The enhancement in the ratio of m/z 39 to 43 can be used to detect the presence of potassium (K^+), since while m/z 39 can have large interferences from the organic fragment C_3H_3^+ , m/z 43 is completely organic⁷¹. In autumn 2018, m/z 39 was highly correlated to m/z 43 (Pearson $R = 0.99$; Fig.S24a), indicating that the fragment was mostly organic. However, an enhancement in m/z 39/43 (over the baseline ratio) was observed during the Diwali period i.e., Nov 7th-8th and on Nov 4th (see temporal trend in Fig.S24b), indicative of the presence of K^+ during these periods. It was likely present during the other periods too given the influence of biomass burning during this season (and the association of K^+ with biomass burning emissions^{24,89-90}) but could not be separated from the organic contribution at m/z 39.

The correlation between m/z 36 (HCl^+) signal and organics mass may indicate the presence of organic chloride⁹¹. However, this method may not be applicable if there are sources that emit organics along with inorganic chloride. In autumn 2018, m/z 36 signal was correlated to organics (Pearson $R = 0.86$; Fig.S25a), but the season experienced sources such as biomass burning (BBOA) and fireworks (sections 3.1 and 3.2) which are known to emit inorganic chloride, e.g. chlorine that may be associated with ammonium and/or potassium from biomass burning, as seen in Nepal during the NAMaSTE study^{92,93}, and other studies around the world⁹⁴⁻⁹⁵ or chlorine associated with potassium from fireworks⁷¹. During autumn 2018, chloride had a very high correlation with BBOA (Pearson $R = 0.95$; Fig.S25b), indicating that it was associated with biomass burning, likely in both organic and inorganic forms, considering the evidence of K^+ and a moderate correlation with NH_4^+ (Pearson $R = 0.75$; Fig.S25b). Ammonia (NH_3) may be emitted from sources other than biomass burning, but can neutralise HCl associated with biomass burning to form particulate ammonium chloride (NH_4Cl)⁹²⁻⁹³. Further, similar to BBOA, average chloride fraction increased with increasing NR-PM_{10} , (see scatter plot in Fig.S26). Furthermore, comparison of normalized weekly moving averages of the PMF factors and inorganic species (Fig.S27; normalized to the average concentration observed in the last two weeks of september) indicates that while all concentrations of all the species increased with the onset of autumn, the increases in chloride and BBOA were most notable (both increased by a factor of 20-30) and correlated with each other. These observations are consistent with the concurrent emissions of organic aerosol and chloride from biomass burning as also observed in other studies⁹⁶⁻⁹⁸.

3.4 Insights from meteorological data

We used non-parametric wind regression analysis to understand the influence of wind speed and direction. We tested the statistical significance of the correlations of average concentrations with meteorological parameters (see scatter plots in Figs.S28-29) by using a significance level of 0.05 (i.e., a p-value <0.05 indicates that the correlation was statistically significant). The p-values and adjusted R^2 for the correlation between different variables are presented in Tables S4-5 and are discussed below.

3.4.1 Wind speed and direction

As shown in the average wind rose (Fig.4a), autumn 2018 observed winds from north west, east and south east. Further, on average, winds from northwest (NW) were relatively faster ($>50\%$ of wind speeds $>3 \text{ ms}^{-1}$). Non-Parametric wind regression (Fig.4b) shows BBOA sources in the NW and S directions. The source in NW is likely agricultural burning from the northwestern states of Punjab and Haryana, considering the spread across NW region at higher wind speeds (regional contributions are expected to increase with faster winds). On average, BBOA concentration decreased with increase in wind speed (Fig.S28a and Table S5), reflective of the contribution from “local” biomass burning sources (e.g. domestic heating), considering that local contributions are expected to decrease with increase in wind speed. This observation is consistent with the BBOA MS which saw contribution from local and aged sources, as discussed in section 3.2.

Figure 4: a) Average wind rose for autumn 2018. The radial values correspond to % frequency. Non-Parametric wind regression of b) BBOA, c) HOA and d) OOA. The radial values correspond to the wind speeds in ms⁻¹

For the HOA factor (Fig.4c), the sources are in E/SE, likely due to proximity of the major roads of Aurobindo Marg and Outer Ring Road (see location in Fig.S30). HOA is also inversely related to wind speed (Fig.S28a), corroborating that it originated from local sources. The OOA concentration is distributed over different speed/wind directions in the NWR plot (Fig.4d), as expected from secondary OA. However, it is slightly inversely correlated to wind speed (Fig.S28a and Table S5). This observation combined with a high oxidation state (section 3.2) is consistent with rapid photochemical processing of emissions contributing to OOA. Among the non organic NR-PM₁ species (see NWR plots in Fig.S31), the NWR plots of Cl⁻ and NH₄⁺ are similar to BBOA, indicative of the association of Cl⁻ with NH₄⁺ and BBOA, in line with our findings in section 3.3. SO₄²⁻ concentration is spread across all directions and

windspeeds, reflective of its regional nature. NO_3^- observed higher concentrations from NW and S at lower wind speeds, reflective of the local nature of NO_x pollution, which is associated with the formation of inorganic and organic nitrate, which constitute the NO_3^- observed in the ACSM.

3.4.2 Temperature, relative humidity, planetary boundary layer height and ventilation coefficient

The concentrations of primary OA PMF factors (BBOA and HOA) decreased with increasing temperature (T), planetary boundary layer height (H) and ventilation coefficient ($\text{VC} = \text{wind speed} \times \text{planetary boundary layer height}$) (Figs.S28 b,d,e and Table S5). The decrease with T is presumably due to evaporation of semivolatile species at higher T; the decrease with H and VC is due to dilution. OOA concentrations did not have a statistically significant correlation with T or H, presumably due to a correlation between photochemistry and higher T and H, which would off-set the effects of partitioning and dilution. With respect to relative humidity (RH), average HOA and BBOA increased up to RH $\sim 75\%$ and decreased afterwards (Fig. S28c). This was likely due to the primary emission peaks coinciding with the time at which average RH was around 75% (Fig.S20) and the influence of precipitation at high RH ($>90\%$; such high RH values were observed mostly during the last two weeks of Sep which marks the end of monsoon season). Amongst the non-organic NR- PM_1 species, average Cl^- , NO_3^- , and NH_4^+ concentrations decreased with increasing T, H and VC (Fig.S29b,d,e and Table S5). SO_4^{2-} did not have a statistically significant correlation with any of these variables (Table S5), reflective of its non-volatile and regional nature. With respect to RH, Cl^- , NO_3^- , and NH_4^+ concentrations increased with increasing RH and decreased at very high RH ($>90\%$) likely due to precipitation (Fig.S29c). SO_4 also increased with increasing RH, albeit to a lesser

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extent than the other species. This is consistent with other recent studies around the world⁹⁸⁻
⁹⁹ that have hypothesised the transformation of gaseous HNO₃ and HCl into aqueous phase
particles at high RH. Another reason could be that high RH usually coincides with low T
which facilitates increased partitioning of NO₃⁻ and Cl⁻ to particle phase.

Thus, temperature and ventilation played an important role in governing the primary OA
species (HOA, BBOA) and the more volatile and local non-organic NR-PM₁ species (Cl⁻, NH₄⁺
and NO₃⁻). Overall, primary emissions and their interplay with meterology played a
dominant role in influencing the air pollution levels during one of the most polluted seasons
in Delhi. Specifically, we observed that while HOA was mostly emitted from local sources
(traffic), BBOA had contributions from regional agricultural burning in addition to local
sources, as opposed to winter, which sees mostly local biomass burning sources. Further,
particulate Cl⁻ also had contributions from biomass burning sources. Diwali period observed
contributions from metals associated with fireworks. Future studies would benefit from
mobile measurements using high resolution instruments in the region to quantity and
apportion the influence of regional versus local sources contributing to BBOA and PM₁
during this season. From a policy perspective, it is evident that a combined effort at state
level (to control local sources) and central level (to control regional sources such as
agricultural burning) are necessary to control air pollution in the city.

Associated Content

Supplementary Information

Details on ACSM calibration and data processing (section S1), PMF analysis (section S2), non-parametric regression (section S3), thirty one figures and five tables supporting the main manuscript.

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TOC graphic (For Table of Contents only)

a)

b)

c)

d)

Oxidizing
capacity

Primary
emissions

NW winds

Rapid
Processing

Moderate
T/PBLH

Regional
transport

Legend for chemical species:

BBOA	HOA	OOA	
Chl	NH ₄	NO ₃	SO ₄